





# Low frequency changes in the carbon cycle during the last 120 kyr and its implications for the reconstruction of atmospheric $\Delta^{14}\text{C}$ and the $^{14}\text{C}$ production rates estimates — a simulation study

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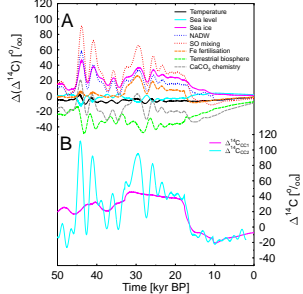
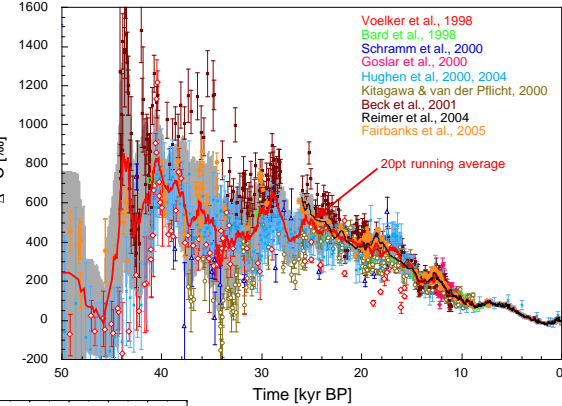
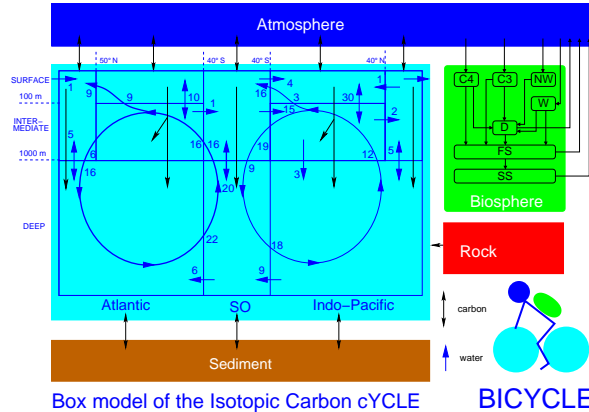
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We use the ocean/atmosphere/biosphere box model of the global carbon cycle BICYCLE (Köhler et al. 2005) to reproduce low frequency changes in atmospheric  $\text{CO}_2$  as seen in Antarctic ice cores during the last glacial cycle (~ 120,000 years) (Köhler et al. 2006). We force the model forward in time by various paleo-climatic records derived from ice and sediment cores. The simulation results of our proposed scenario match a compiled  $\text{CO}_2$  record from various ice cores with high accuracy ( $r^2 = 0.89$ ). The processes that contribute most to the glacial/interglacial changes in  $\text{CO}_2$  are variations in the sedimentation and dissolution rates of  $\text{CaCO}_3$ , ocean circulation, ocean temperature and glacial iron fertilization of the marine biota in the Southern Ocean. The BICYCLE model includes also calculations for the carbon isotopes  $^{13}\text{C}$  and  $^{14}\text{C}$  and we assess what changes in atmospheric  $\Delta^{14}\text{C}$  might be based on variations in the carbon cycle. Our results suggest that during the last glacial cycle in general less than 120‰ of the increased atmospheric  $\Delta^{14}\text{C}$  are based on variations in the carbon cycle, while the largest part of the variations has to be explained by changing  $^{14}\text{C}$  production rates. Processes acting on the global carbon cycle that increase glacial  $\Delta^{14}\text{C}$  are a restricted glacial gas exchange between the atmosphere and the surface ocean through sea ice coverage, a reduced glacial ocean circulation, and the enrichment of DIC with  $^{14}\text{C}$  in the surface waters through isotopic fractionation during higher glacial marine export production caused by iron fertilization. From the available  $\Delta^{14}\text{C}$  data covering the last 50,000 years and our carbon cycle-based simulation results we can infer changes in the  $^{14}\text{C}$  production rates, which are then compared with two other estimates based on  $^{10}\text{Be}$  and geomagnetic field reconstruction. The agreements and discrepancies between these three independent approaches to estimate the  $^{14}\text{C}$  production rates are discussed and highlight the limitations and possible uncertainties in all three approaches.

**Keywords:** carbon cycle,  $^{14}\text{C}$  cycle,  $^{14}\text{C}$  production rates, glacial/interglacial, modeling, box model, radionuclides

## Model — $\Delta^{14}\text{C}$ data — Results

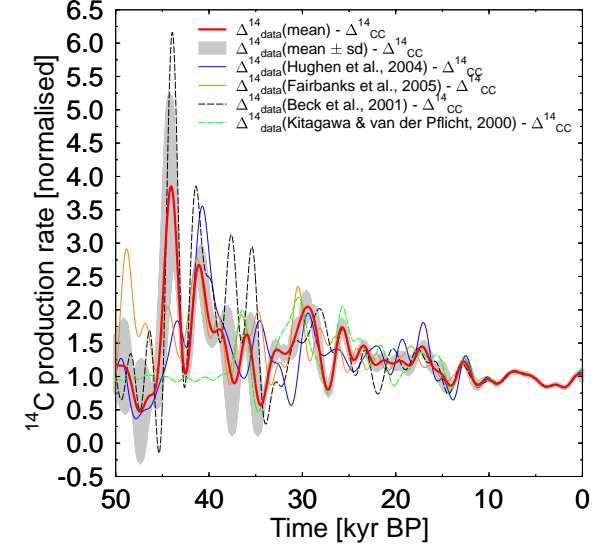


Top: Data based reconstruction of  $\Delta^{14}\text{C}$ .

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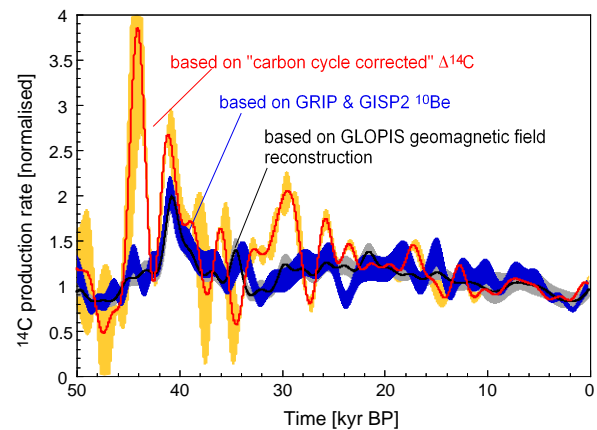
Top: Contribution of different processes and overall carbon-cycle based changes to  $\Delta^{14}\text{C}$  with variable  $^{14}\text{C}$  production rates.

## Variable $^{14}\text{C}$ production rate



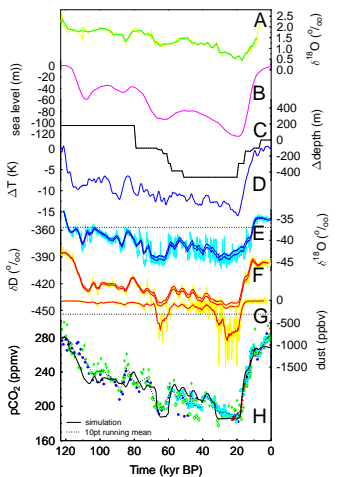
Carbon cycle-based reconstruction of  $^{14}\text{C}$  production rate based on different  $\Delta^{14}\text{C}$  data sets.

## Comparing three approaches to estimate $^{14}\text{C}$ production rates

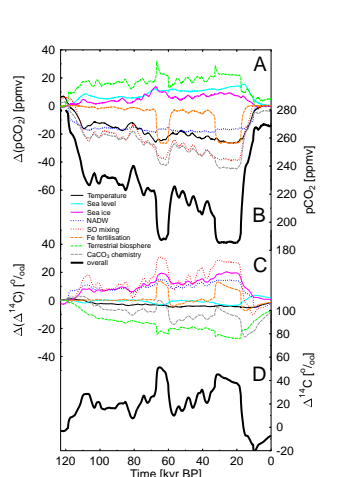


$^{10}\text{Be}$  from Muscheler et al. (2005) and GLOPIS from Laj et al. (2005).

## Forcing



## Simulation Results



Left: Different data sets forcing the model (A: equatorial SST proxy; B: sea level; C: lysocline; D: northern hemisphere temperature; E: North Atlantic SST proxy; F: Southern Ocean SST proxy; G: dust input in Southern Ocean; H: Data and simulation results for  $\text{pCO}_2$ .  
 Right: Simulation results. A: Process contribution to  $\Delta(\text{pCO}_2)$ ; B:  $\text{pCO}_2$ ; C: Process contribution to  $\Delta(\Delta^{14}\text{C})$ ; D:  $\Delta^{14}\text{C}$  with constant  $^{14}\text{C}$  production rate.

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