

# Thermochronological Research in Northern Victoria Land (Antarctica): a Key to the pre-Disintegration Palaeogeography of Panthalassian Gondwana

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**Abstract:** This paper presents the overview of the apatite fission track (AFT) data set of northern Victoria Land and juxtaposed SE Australia, and the underlying geological and geomorphological processes since late Paleozoic times. It focuses on thermochronological data produced during the last two decades, new interpretation strategies, and the combined use of thermochronological data and complementary geological information. The regional AFT pattern and stratigraphic age information require the existence of a late Paleozoic - Mesozoic sedimentary basin between northern Victoria Land and SE Australia. Basin formation resulted from long-lasting N-S oblique extension that also triggered the ~180 Ma Ferrar magmatism, the rifting of the Ross Sea, and eventually continental breakup between Antarctica and Australia. The locus of breakup is probably controlled by basin geometry and depth.

**Zusammenfassung:** Dieser Artikel präsentiert einen Überblick über den Datensatz von Apatitspaltspurenaltern (AFT) von Nordviktoraland und dem gegenüberliegenden Südosten Australiens sowie über die zu Grunde liegenden geologischen und geomorphologischen Prozesse seit dem Paläozoikum. Er fokussiert auf thermochronologische Daten der letzten beiden Jahrzehnte, neue Interpretationsstrategien sowie die Kombination thermochronologischer Daten mit komplementären geologischen Informationen. Der Vergleich regionaler AFT-Alter und stratigraphischer Alter erfordert die Existenz eines spät-paläozoisch-mesozoischen Beckens zwischen Nordviktoraland und SE-Australien. Dessen Bildung ist das Ergebnis lang anhaltender schräger N-S-Extension, die auch für den ~180 Ma Ferrar-Magmatismus, das Rossmeer-Rifting und schließlich den Kontinentalzerfall zwischen der Antarktis und Australien verantwortlich ist. Die Anlage der entsprechenden Bruchstelle ist vermutlich durch Beckengeometrie und Beckentiefe kontrolliert.

## INTRODUCTION

Compared to former supercontinents, modern Gondwana reconstructions tend to be reasonably well constrained. They mainly rely on fits of isobaths, positions of palaeomagnetic poles or comparisons of structure and age of the metamorphic basement and Palaeozoic to Mesozoic sedimentary strata including fossil content, respectively, of the juxtaposed Gondwana fragments. Nevertheless, the resolution of the pre-breakup architecture of some regions within Gondwana remains poor, especially of Antarctic regions now covered by ice and/or lacking any attributable sedimentary record. A typical example for this dilemma is the correlation of the sheared margins of the Ross Sea sector of Antarctica and SE Australia (Fig. 1).

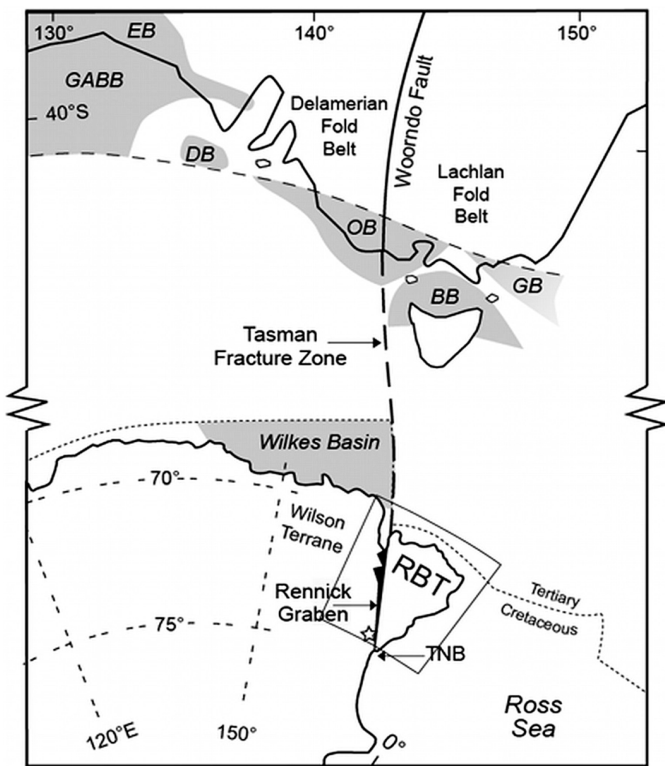
While the coast between South Australia and Victoria is easily accessible, and abundant sedimentary strata are preserved both onshore and on a broad continental shelf, only a few outcrops are exposed at the juxtaposed coast between Terre Adélie and northern Victoria Land, and the shelf there is not accessible. The basement of Terre Adélie and northern Victoria Land consists of the Precambrian East Antarctic Craton onto which three major tectono-metamorphic terranes (Wilson, Bowers, and Robertson Bay) were accreted during the Early Paleozoic Ross Orogeny (Fig. 2). It comprises Neoproterozoic to early Paleozoic gneisses and low-grade metasedimentary rocks intruded by the ~500 Ma Granite Harbour and the ~350 Ma Admiralty Intrusive suites (e.g., BORG et al. 1987). This basement is overlain only locally by remnants of a so-called Gondwana terrestrial sequence made of Permian to Early Jurassic clastic deposits (Beacon Supergroup), and mafic Ferrar volcanic rocks and dykes that emplaced at ~180 Ma (e.g., COLLINSON et al. 1986, BARRETT 1991, HEIMANN et al. 1994). Neogene volcanic rocks and sediments are merely preserved as patches in the immediate vicinity of the Ross Sea margin.

Due to these limitations, early correlations of Antarctica and SE Australia largely depended on proxies for properties of crustal units, such as paleomagnetism, geochronology, magmatic and metamorphic petrology, structural geology (e.g., STUMP et al. 1986, MILLER et al. 2002), whereas sedimentary and stratigraphic information is restricted to the Australian side (e.g., MUTTER et al. 1985, WILLCOX & STAGG 1990, STAGG & WILLCOX 1992, BRYAN et al. 1997).

In this situation, the Antarctic Ross Sea region became one of the first areas where an alternative dating technique, apatite fission track (AFT) analysis, was applied to solve this dilemma. AFT thermochronology is a radiometric method based on the accumulation of damage trails in the mineral apatite due to spontaneous nuclear fission of <sup>238</sup>U. Fission tracks are produced at a constant rate through time, and so the number (density) of tracks can be used to estimate the time since track accumulation began, i.e. the AFT age. Fission tracks remain preserved in apatites below a temperature of 110-125 °C. They experience some length reduction (annealing) within the temperature range of 60-110 °C ("Partial Annealing Zone", e.g., WAGNER & VAN DEN HAUTE 1992). AFT analysis produces two parameters: an AFT age (resulting from the track density), and information about the style of cooling and the maximum paleotemperature (from mean track length and standard deviation). Moreover, etch pit diameters

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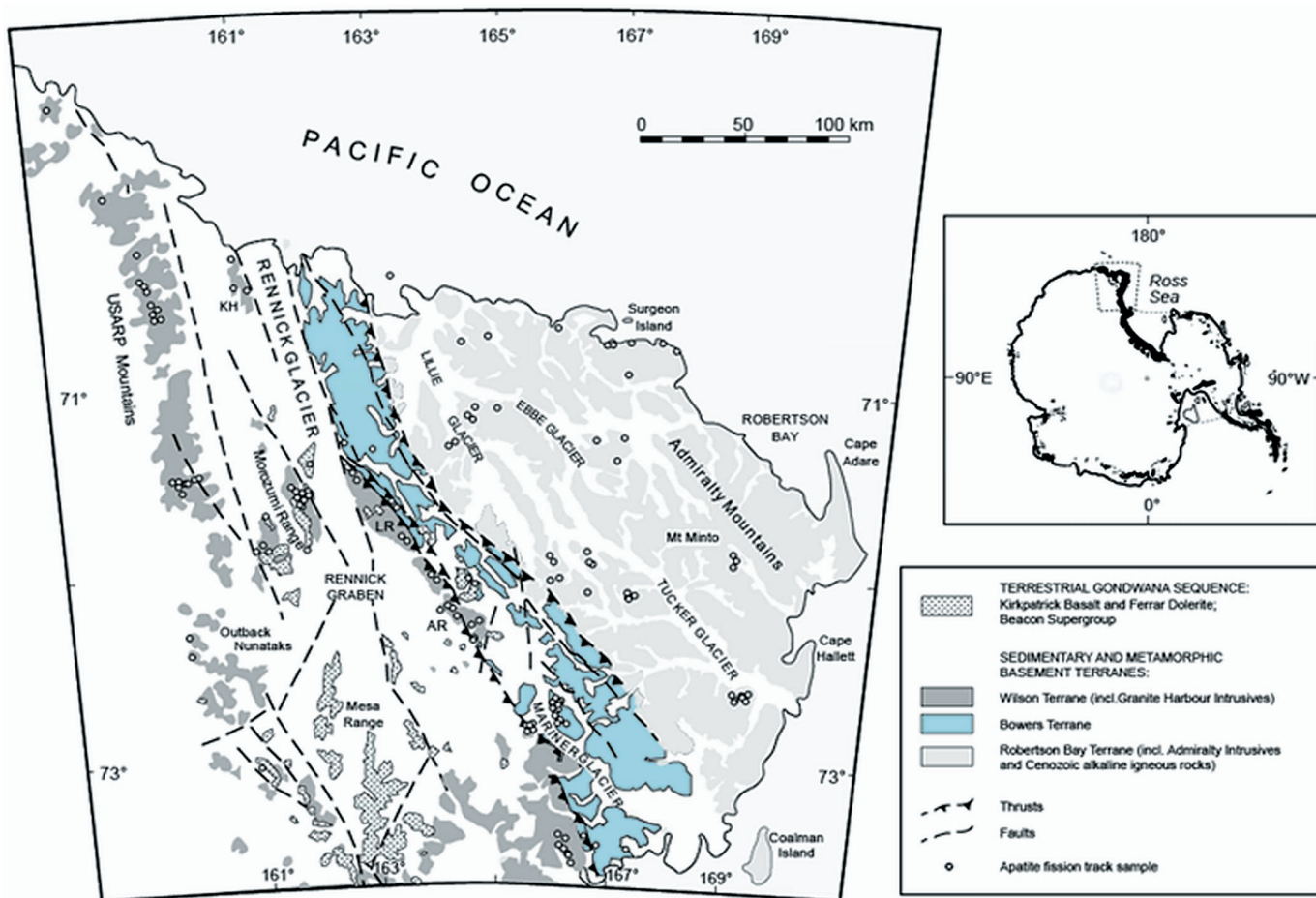


are measured as a proxy for the chemical composition of the apatites (BURTNER et al. 1994).

Qualitative interpretation and thermal history modelling of AFT data are then applied to derive constraints on timing and amount of cooling. Cooling of rocks to the temperatures of the AFT system usually refers to exhumation or postmagmatic thermal relaxation. At present, AFT thermochronology is the most commonly used tool to delineate exhumation processes and long-term landscape evolution of a wide field of geological environments, ranging from contractional to extensional settings and "stable" cratonic interiors. For detailed overviews on apatite thermochronology and its application to geological

**Fig. 1:** Sketch map showing the juxtaposed coastal segments between Terre Adélie and northern Victoria Land of Antarctica and southeastern Australia. Basins: BB = Bass Basin; EB = Eucla Basin; DB = Duntroon Basin; GABB = Great Australian Bight Basin; GB = Gippsland Basin; OB = Otway Basin. RBT = Robertson Bay Terrane; TNB = Terra Nova Bay. The star marks the location of the Eisenhower and Deep Freeze ranges, the frame indicates the contour of Figure 2.

**Abb. 1:** Übersichtskarte mit den gegenüberliegenden Küstenabschnitten zwischen Terre Adélie und Nordviktoraland der Antarktis und dem südöstlichen Australien. Becken: BB = Bass-Becken, EB = Eucla-Becken, DB = Duntroon-Becken, GABB = Great Australian Bight-Becken, GB = Gippsland-Becken, OB = Otway-Becken. RBT = Robertson Bay Terrane; TNB = Terra Nova Bay. Der Stern markiert die Lage der Eisenhower Range und der Deep Freeze Range, der Rahmen zeigt den Umriss der Abbildung 2.



**Fig. 2:** Geological sketch of northern Victoria Land showing sample locations of the apatite fission track studies of FITZGERALD & GLEADOW (1988), BALESTRIERI et al. (1994b, 1997, 1999), LISKER (1996 unpubl. data), SCHÄFER (1998), BALESTRIERI & BIGAZZI (2001), ROSSETTI et al. (2003), and LISKER et al. (2006).

**Abb. 2:** Geologische Übersichtskarte von Nordviktoraland mit den Probenahmeorten für die Spaltspuruntersuchungen von FITZGERALD & GLEADOW (1988), BALESTRIERI et al. (1994b, 1997, 1999), LISKER (1996 unpubl. Daten), SCHÄFER (1998), BALESTRIERI & BIGAZZI (2001), ROSSETTI et al. (2003) und LISKER et al. (2006).

problems we refer to WAGNER & VAN DEN HAUTE (1992), GALLAGHER et al. (1998), GLEADOW et al. (2002a), REINERS & EHLERS (2005), and LISKER et al. (2009).

#### EARLY THERMOCHRONOLOGICAL STUDIES IN ANTARCTICA AND AUSTRALIA

Southeastern Australia and the formerly linked portion of Pacific Antarctica belonged to the first regions where thermochronological techniques were applied and represent key areas for the development of the AFT method and basic principles of the interpretation of thermochronological data. The first geological applications in Australia dated cooling processes of rocks from coastal Victoria and were related to rifting and passive margin evolution (GLEADOW & LOVERING 1978a,b, GLEADOW & DUDDY 1981, MOORE et al. 1986). Methodological aspects of these studies referred to the importance of fission track length distributions for estimating paleotemperatures, and to the applicability of the method to basin research and hydrocarbon exploration. Early studies in Antarctica concentrated on the Transantarctic Mountains (GLEADOW et al. 1984, FITZGERALD 1986, GLEADOW & FITZGERALD 1987, FITZGERALD & GLEADOW 1988, FITZGERALD & STUMP 1991, FITZGERALD 1992, 1994, cf. reviews of BALESTRIERI et al. 1994a and FITZGERALD 2002) and developed diagnostic tools to determine timing and amount of "uplift"/ exhumation (e.g., the "break in slope" representing a fossil Partial Annealing Zone: GLEADOW & FITZGERALD 1987, FITZGERALD & GLEADOW 1990). The AFT data set of the Ross Sea region was interpreted in terms of three episodes of "uplift" – in modern terminology: exhumation stages – related to regional tectonic events (e.g., FITZGERALD 2002):

- (1) the initial breakup between Australia and Antarctica in the Early Cretaceous,
  - (2) the main phase of extension between East and West Antarctica in the Late Cretaceous, and
  - (3) the propagation southward of seafloor spreading from the Adare Trough into continental crust underlying the western Ross Sea in the early Cenozoic.
- This latter event likely acted as the trigger for the flexural uplift of East Antarctic lithosphere to form the Transantarctic Mountains.

The AFT data also provide a main base for various uplift scenarios that can be divided into four general groups: thermally driven uplift, mechanically driven uplift (crustal extension and flexure, a combination of these two (e.g.), or topographic reversal due to the collapse of a West Antarctic plateau (SMITH & DREWRY 1984, FITZGERALD et al. 1986, SALVINI et al. 1997, ten BRINK et al. 1997, LAWRENCE et al. 2006, FACCENA et al. 2008, BIALAS et al. 2007, VAN WIJK et al. 2008). Much less attention has been given to the Pacific continental margin that evolved perpendicular to the West Antarctic Rift System. Early age data from this margin were not obtained from independent research projects but collected within studies focusing on the northern segments of the Transantarctic Mountains.

#### FISSION TRACK COMPILATIONS OF THE ANTARCTIC - AUSTRALIAN MARGIN

The first reviews of AFT data in the context of the passive margin evolution of Antarctica and Australia were published by STUMP et al. (1990) and FOSTER & GLEADOW (1992, 1993). These studies derived the regional "uplift" history of this part of Gondwana from AFT data produced by GLEADOW & LOVERING (1978a, b), MOORE et al. (1986), and DUMITRU et al. (1991; all Victoria), and by GLEADOW & FITZGERALD (1987) and FITZGERALD & GLEADOW (1988; both Victoria Land). Both juxtaposed margin segments show a consistent pattern of old AFT ages (up to ~400 Ma) in their western terranes (Delamerian Fold Belt/Australia, Wilson Terrane/ northern Victoria Land) while the ages of the eastern terranes (Lachlan Fold Belt, Robertson Bay Terrane) are usually <100 Ma. This age pattern confirms the match of the Gondwana terranes as proposed earlier by STUMP et al. (1986) on the base of stratigraphic and lithological similarities, and highlights the importance of a regional tectonic lineament consisting of Woorndoo-Sorrel Fault Zone – Tasman Fracture Zone – Leap Year Fault (Rennick Graben) (FOSTER & GLEADOW 1992) (Fig. 1).

The compilations also propose a common thermal history of all terranes from the Devonian through to the end of the Mesozoic that comprised very little burial or exhumation. A thermal reset was recognized for the ~180 Ma Ferrar magmatism in northern Victoria Land, but neither the SE Australia nor northern Victoria Land data show any clear influence of rifting and breakup in the late Cretaceous. Subsequent to the breakup of Australia and Antarctica, the thermal and tectonic histories of both margins evolved independently along differing paths. With respect to present-day geomorphic differences, STUMP et al. (1990) suggested that during extension northern Victoria Land was flanked by two upper plate margins, whereas southeastern Australia was flanked by an upper plate and a lower plate margin. FOSTER & GLEADOW (1992, 1993) particularly focused on the lithospheric boundaries across the margins that were supposedly reactivated as transfer faults during Mesozoic rifting and Gondwana fragmentation. They now divide crustal segments of different rheological behaviour, and with varying amounts of uplift.

A review of LISKER (2002) based on a much larger body of published AFT data from the Transantarctic Mountains (STUMP & FITZGERALD 1992, BALESTRIERI et al. 1994b, 1997, FITZGERALD et al. 1996, FITZGERALD & STUMP 1997) including three comprehensive data sets from PhD theses (LISKER 1996, MILLER 1997, SCHÄFER 1998). It provided a better resolution of thermal histories, and extended the Antarctic AFT data coverage further East towards Oates Land. All AFT compilations of the juxtaposed Antarctic and Australian margins stress the similarity of the AFT data pattern on both continents, and agree in the correlation of terranes and tectonic lineaments as well as structural control of exhumation. In addition, LISKER (2002) refers to the influence of plate rotation on exhumation and uplift.

The review papers and all underlying studies consider the heterogeneous passive/ transform margins of northern Victoria Land and SE Australia as the result of long-lasting rotation, extension and rifting that started in the Jurassic, with sea floor production commencing in the Early Cretaceous. They suggest

that a change in stress pattern triggered a discrete sequence of two events comprising late Mesozoic continental rifting and separation, and Cenozoic West Antarctic Rifting. In addition, all authors explicitly or implicitly consider the Transantarctic Mountains as a long-lasting, stable mountain chain/highland that might have brought into existence as early as ~180 Ma during the Ferrar event. It was argued to then have been uplifted stepwisely at least since Early Cretaceous times. However, this concept produces a number of problems and open questions concerning data interpretation and the relationship between AFT data and other geological and geomorphological evidence.

Firstly, the postulated landscape evolution model comprising three episodes of rapid uplift/exhumation based on qualitative interpretation cannot be verified by thermal history modelling of the AFT data because the track lengths are too short in most samples. These models further conflict with the rather homogeneous distribution of a series of reference horizons, such as erosion surfaces, sedimentary and volcanoclastic strata, volcanic flows, and with various thermal indicators (cf. LISKER & LÄUFER 2007).

Secondly, uplift of the different segments of the Transantarctic Mountains was not recorded simultaneously and according to a regular trend along the mountain chain, but instead appears diachronous and without a recognizable spatial pattern (cf. FITZGERALD 2002). Northern Victoria Land, constituting the northernmost segment of the Transantarctic Mountains in the Ross Sea region lacks any consistent interference of breakup-related exhumation and exhumation induced by rifting of the West Antarctic Rift System. Moreover, combined thermochronological and structural data indicate a repeated swap of the regional stress field into perpendicular directions between Jurassic and Paleogene.

Of particular importance with respect to Gondwana breakup and subsequent transform/passive margin evolution is the contrast between Cretaceous deposition on the Australian margin including a thick shelf sequence and supposed contemporaneous exhumation of the Antarctic margin in spite common AFT patterns across both margins.

These inconsistencies between AFT data and complementary geological information requires a state of the art reconstruction of the breakup processes between Antarctica and Australia that has to rely on four parallel avenues: new AFT data, new techniques, a focus on isotherm patterns, and the intimate link to various geoscience disciplines.

## AFT WORK OF THE LAST DECADE

### (i) *New AFT data*

Since the publication of the main reviews of AFT data from the Australian and Antarctic margin in the late 1990's, a decade of further study has significantly expanded the AFT dataset. AFT studies between northern Victoria Land and Terre Adélie (Antarctica) were conducted by BALESTRIERI & FIORETTI (1998), BALESTRIERI et al. (1999), BALESTRIERI & BIGAZZI (2001), FITZGERALD (2001), ROSSETTI et al. (2003), LISKER & OLESCH (2003, 2004), LISKER et al. (2006), STORTI et al. (2008), MILLER et al. (2010), and ZATTIN et al. (2010).

When added to the existing data, these studies have contributed to form an overall data set of more than 500 AFT ages and associated proxies.

The vast AFT data base from the Australian continent comprises more than 3000 records, with the majority of them obtained from SE Australia (Victoria, New South Wales, and Tasmania) (GLEADOW et al. 2002b, KOHN et al. 2002). Many of these data originated during hydrocarbon exploration in the onshore and offshore basins (HILL et al. 1995, O'SULLIVAN et al. 1995, 1996, 1999, 2000a, b, c, MITCHELL 1997, KOHN et al. 1999, GIBSON & STÜWE 2000, MITCHELL et al. 2002, TINGATE & DUDDY 2002, GREEN et al. 2004, and WEBER et al. 2004).

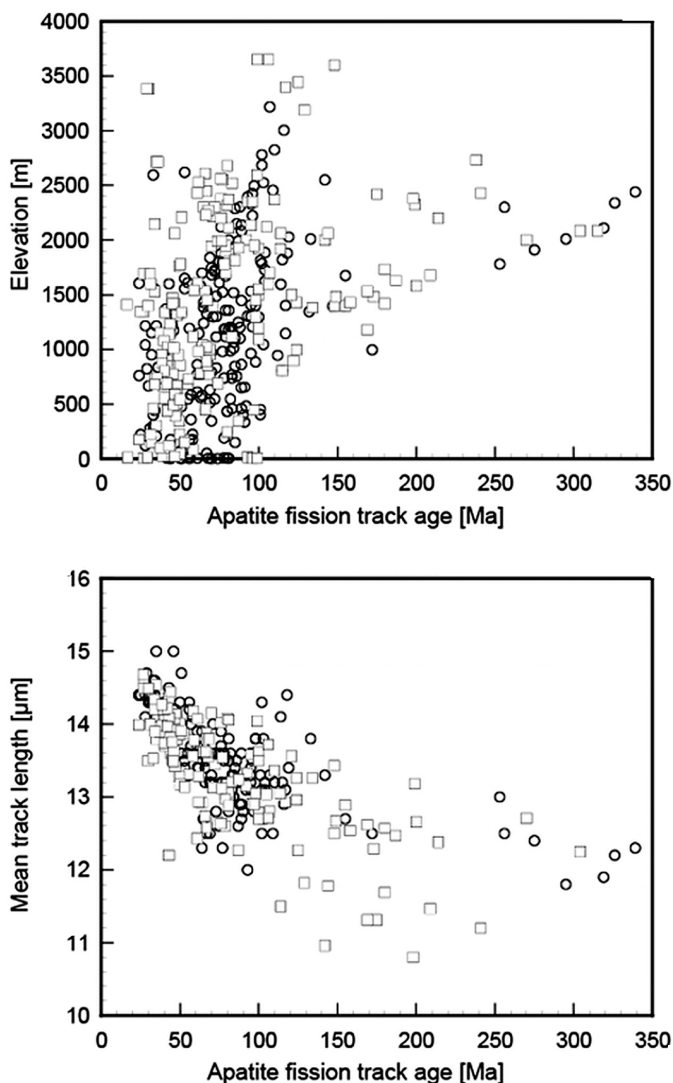
More important than the addition of new apatite ages are the extension of the study areas away from the dominant rift/margin structures in the east where thermal history reconstruction is superimposed and "blurred" by younger tectonic processes, as well as an improved thermal resolution due to the addition of a large quantity of annealing proxies (AFT length data, etch pit diameters/ $D_{par}$ ). Some key target areas containing high-resolution vertical AFT profiles (e.g. from escarpments or boreholes) and/or horizontal transects do not only provide excellent temperature constraints at various time intervals but also allow the calculation of paleogeothermal gradients at the time of maximum burial (e.g., O'SULLIVAN et al. 2000a, MITCHELL et al. 2002, GREEN et al. 2004, WEBER et al. 2004, LISKER et al. 2006).

Age spectra and regional distribution of the AFT data generated during the last decade very much resemble the ones of the former studies, and constrain a systematic bimodal pattern consisting of predominantly Late Cretaceous to early Cenozoic AFT ages coupled with relatively long track lengths and a broad range of older ages linked with usually short track lengths. The majority of samples, including those with AFT ages of ~50 Ma have mean track lengths shortened to below 14  $\mu\text{m}$  (Fig. 3). In general, the Cenozoic ages dominate the eastern coastal regions of both continental fragments, and increase towards west (Fig. 4). In northern Victoria Land, this trend usually correlates with the geomorphological contrast between Alpine topography at the Ross Sea and the plateau bound by steep escarpments towards the west, whereas the youngest AFT ages in SE Australia are confined to the Great Escarpment (e.g., GLEADOW et al. 2002b, KOHN et al. 2002; see below). LISKER (2002) identified the line Rennick Graben – Tasman Fracture Zone – Woorndoo-Sorrel Fault Zone as a key structure dividing the age pattern (Fig. 4).

In contrast, there is no distinctive N–S trend perpendicular to the Antarctic coast as typical for passive margin settings. This contrasts with a trend of young ages/large amounts of exhumation at the coast and increasing ages/decreasing exhumation towards the interior in the southeast of Australia (e.g., KOHN et al. 2002).

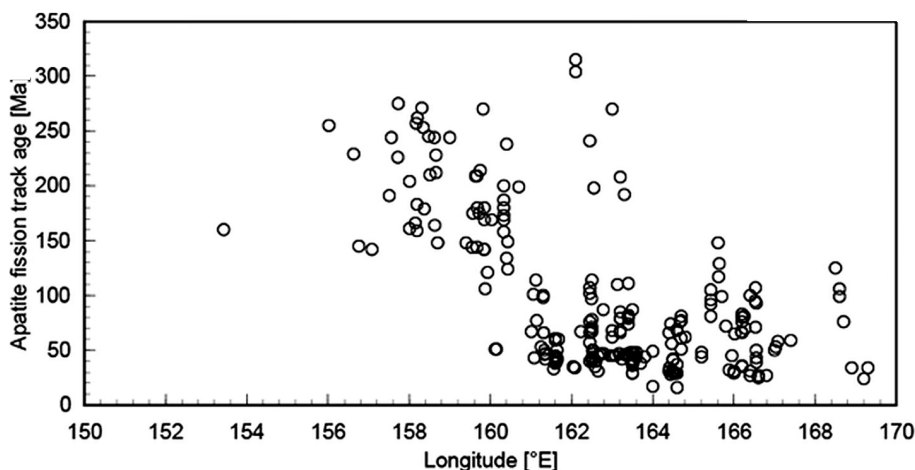
### (ii) *(U-Th)/He thermochronology: a new, supplementary method*

The AFT technique represents a unique tool to investigate the thermal history of rocks within the temperature range of 60 - 110 °C, usually referring to exhumation processes within



**Fig. 3:** Plots showing the relationship between apatite fission track age and elevation (top) and the distance to the “transantarctic” Ross Sea margin (bottom), respectively. Circles = data from northern Victoria Land (Fig. 2) and the Terra Nova Bay; squares = data of the central and southern Transantarctic Mountains. Note the uniform trend of both data sets.

**Abb. 3:** Grafik mit dem Verhältnis zwischen Apatit-Spaltspuralter und Höhe (oben) und der Entfernung zum “transantarktischen” Rossmeerrand (unten). Kreise = Daten aus Nordviktoraland (Abb. 2) und der Terra Nova Bay; Quadrate = Daten aus dem zentralen und südlichen Transantarktischen Gebirge. Man beachte den einheitlichen Trend beider Datensätze.



**Fig. 4:** Diagram showing the trend of apatite fission track (AFT) ages along the coast of northern Victoria Land. Note the sudden shift of the AFT ages across the western master fault of the Rennick Graben at 163°–164°E. Ages from the cratonic interior west of the Morozumi Range are exclusively >100 Ma while the majority of the terrane sample ages near the Ross Sea are <100 Ma.

**Abb. 4:** Diagramm zur Verdeutlichung des Trends der Apatit-Spaltspur (AFT)-Alter entlang der Küste Nordviktoralands. Man beachte den plötzlichen Sprung der AFT-Alter über die westliche Haupttrandstörung des Rennick-Grabens bei 163–164°E. Alter aus dem Kratonbereich westlich der Morozumi Range sind ausschließlich älter als 100 Ma, während die Mehrheit der Alter der aus den Terranes nahe des Rossmeeres stammenden Proben jünger als 100 Ma sind.

crustal depths of 2-5 km. However, the restriction to these temperatures does not provide access to the youngest cooling/exhumation phases of the uppermost crust, especially for areas with old ages lacking clear signatures. Additional data of higher thermal sensitivity can be acquired using (U-Th)/He thermochronology. This methodology relies on the radioactive decay of U and Th to  ${}^4\text{He}$  to determine the time since an apatite crystal cooled through the temperature interval 40-85 °C (e.g., WOLF et al. 1998). It is applied to investigate the vertical throws along faults, the rates of continental erosion, the formation of topographic relief and climate variation (e.g., EHLERS & FARLEY 2003). The first (U-Th)/He data from the Transantarctic Mountains were published by FITZGERALD et al. (2006). This study dated apatite aliquots from vertical sample profiles in southern Victoria Land that were analyzed previously for fission track data. The (U-Th)/He ages were usually 10-20 Ma younger than corresponding AFT ages (43-92 Ma). Similarly, HOUSE et al. (1999, 2002) produced apatite (U-Th)/He ages of surface samples from the Otway Basin (SE Australia; ~70 Ma) that are substantially younger than corresponding AFT ages (~110 Ma). PERSANO et al. (2002, 2006) obtained similar relationships from geomorphological and exhumation studies at the Great Escarpment.

### (iii) Isotherm patterns and recognition of nonlinear cooling

Unlike many geochronological ages, AFT ages usually cannot be used as direct time constraints for immediate dating of discrete geological events. Accordingly, the sole compilation of AFT age data is only of limited use for reconstructing exhumation patterns and tectonic processes. Instead, the advantage of thermochronological methods is the potential to qualitatively estimate and quantitatively model temperature constraints at various times. This potential has been demonstrated by GLEADOW et al. (2002b) and KOHN et al. (2002) for Australia. Moreover, O’SULLIVAN et al. (1995, 1996, 2000a, b,c) applied the concept of nonlinear cooling to the Lachland Fold Belt, the Bathurst region, and the Bassian Rise.

### (iv) Link to complementary disciplines

A reliable interpretation of thermochronological data fundamentally depends on quality and substance of supplementary information from various associated geoscience disciplines.

This includes both a compilation of common features and the definition of critical differences. Beyond thermochronological data, valuable information is available both from academic research and hydrocarbon exploration on the Australian shelf in various disciplines, such as reference horizons, paleotemperatures, stratigraphic patterns, kinematic indicators, geophysical data, geochronological ages, landsat TM analysis, and others. The more pertinent observational and analytical datasets are reference horizons.

### Reference horizons

The potential of such horizons was recognized in the early review papers (e.g., STUMP et al. 1990) both as an independent indicator for burial and exhumation ("uplift") as well as link between southeastern Australia with northern Victoria Land. Very useful are paleosurface markers, most notably unconformities or unconformities, as key indications for the application of simple linear or non-linear cooling/exhumation scenarios. The crucial Antarctic marker horizons comprise:

- (1) The Paleozoic pre-Beacon ("Kukri") Surface on which the strata of the Beacon Supergroup were deposited between the Devonian (central Transantarctic Mountains) to Permian (northern Victoria Land) and the Early Jurassic (e.g., BARRETT 1991, ISBELL 1999).
- (2) the ~350 Ma rhyolitic equivalents of the Admiralty Group which were extruded directly over the deformed pre-Ordovician basement (e.g., FIORETTI et al. 1997).
- (3) The ~180 Ma Ferrar volcanic rocks that conformably overly, or intrude at shallow depths the Beacon Supergroup.
- (4) The various small-scale Cenozoic volcanics and shallow intrusions (for example, Meander, Malta, Hallett magmatics; e.g., ROCCHI et al. 2002).

Some of these features are only preserved as relics and/or their age is poorly resolved or diachronous (Kukri surface, Carboniferous rhyolites, the lower Beacon strata), while others are widely distributed and provide well-defined time constraints (Ferrar volcanics).

Of these palaeosurface markers, the mafic Ferrar Dolerite suite is potentially the most controversial. It is often considered to consist of sills intruded at various depths with only a minor effusive component called Kirkpatrick basalt (e.g., ELLIOT 2000). However, systematic research during the last decade established the predominantly subaerial or very shallow nature of the Ferrar rocks. This is based on:

- (1) syn-Ferrar pyroclastic and partially fossil-bearing siliciclastic sedimentary sequences (e.g., ELLIOT 1996, 2000, SCHÖNER et al. 2007, 2011 this vol.),
- (2) the presence of pillow lavas,
- (3) phreatomagmatic structures and diatremes of local hydro-magmatic explosive events,
- (4) by the content of vesicles and the chilled contacts of sediment suspensions in sills, and
- (5) by the plastic deformation of Jurassic sediments by Ferrar apophyses (e.g., VIERECK-GÖTTE et al. 2007, ELLIOT & FLEMING 2008).

Equivalent marker horizons are available from the Australian side, with the Tasmanian Jurassic dolerites and the Cretaceous Whitsunday volcanics (e.g., BRYAN et al. 1997) being the most relevant ones.

### PALEOTEMPERATURE COMPILATION

Paleotemperature constraints for northern Victoria Land were derived from the Gondwana terrestrial sequence and basement rocks. These include for example diagenetic features in Beacon Supergroup, remagnetisation within low-grade metamorphic rocks, the disturbance of Rb-Sr, K-Ar and Ar/Ar systems of Ferrar rocks, secondary mineral paragenesis within Ferrar rocks, and epidote on brittle fault planes in Ross and post-Ross rocks (e.g., KREUZER et al. 1981, DELISLE & FROMM 1984, 1989, SCHMIERER & BURMESTER 1986, FLEMING et al. 1992, 1993, 1999, FAURE & MENSING 1993, HORNIG 1993, MENSING & FAURE 1996, MOLZAHN et al. 1999, BALLANCE & WATTERS 2002, BERNET & GAUPP 2005). These palaeotemperature constraints refer to maximum post-Jurassic temperatures between <60 and ~350 °C for the now exposed surface rocks. MOLZAHN et al. (1999) relate most of these temperatures to two thermal pulses during the Early and Late Cretaceous. Post-orogenic paleotemperatures of similar magnitude, but unconstrained timing, are also derived from Paleozoic low-grade metamorphic rocks from northern Victoria Land. They include mineral assemblage and metamorphic zonation of low-pressure pelite and calc-silicate rocks, fluid inclusions within Granite Harbour Intrusives, Admiralty Intrusives and metamorphic rocks, quartz recovery, illite crystallinity, and conodont colour alteration (BUGGISCH & KLEINSCHMIDT 1991, CRAW et al. 1992, FADDA et al. 1994, CRAW & COOK 1995, FREZZOTTI et al. 1997, ROSSETTI et al. 2006).

Various studies from SE Australia similarly report predominantly Late Cretaceous maximum paleotemperatures between 110 and 250 °C. Thermal constraints were derived from zeolite assemblages, vitrinite reflectance, fluid inclusion, and geomagnetic data (e.g., SUTHERLAND 1977, MIDDLETON & SCHMIDT 1982, KENNARD et al. 1999, GEORGE et al. 2004).

### STRATIGRAPHY

The terrestrial Gondwana sequence preserved near the Antarctic coast consists of limited and relatively thin ( $\leq 300$  m) exposures of clastic and volcanoclastic sediments of Permian to Early Jurassic age. These crop out in the vicinity of the Rennick Graben (e.g., COLLINSON et al. 1986).

In contrast, SE Australia still contains a whole series of late Mesozoic basins aligned along the southern coast (Fig. 1). Of these, the Eucla/Great Australian Bight, Duntroon, Otway, Bass, and Gippsland basins were directly linked with the now juxtaposed coast between northern Victoria Land and Terre Adélie (cf., MUTTER et al. 1985). The thickness of the predominantly Cretaceous sedimentary sequences within these coastal basins regularly exceeds 5 km, while the adjacent shelf is overlain by up to 15 km sedimentary strata of mainly Jurassic and Early Cretaceous age. These shelf sediments form a 4-5 km thick blanket on continental basement still 120 km offshore.

### STRUCTURAL DATA

Brittle kinematic indications in basement and cover rocks of Victoria Land have been studied intensely during the last

decade (e.g., ROSSETTI et al. 2002, 2003, 2006, LÄUFER et al. 2003, STORTI et al. 2006, KLEINSCHMIDT & LÄUFER 2006, LÄUFER et al. 2011 this vol.). These data constrain the Cenozoic tectonic evolution and suggest this region experienced dextral shear and extensional to transtensional tectonics during the formation of the Ross Sea (e.g., SALVINI et al. 1997). The strike-slip faults cutting through the continental crust of northern Victoria Land are interpreted to be the direct prolongation of the intra-oceanic fracture zone arrays between Australia and Antarctica (e.g., the Tasman Fracture Zone). Thermochronological data provide absolute time constraints to this deformation. For example, the oblique rifting of the West Antarctic Rift System since ~50 Ma was triggered by the transfer of lithospheric stress and right-lateral shear from the Southern Pacific Ocean into the Antarctic crust of northern Victoria Land (e.g., ROSSETTI et al. 2006, STORTI et al. 2008).

Similar transtensional tectonics are argued to have occurred in Australia. MILLER et al. (2002) for example report the development of Cretaceous extensional to transtensional faults continuing offshore to define the oceanic transform faults between Australia and Antarctica. The positioning of these was largely controlled by pre-existing structural lineaments, which date to the Delamerian and Lachlan orogenic events (MILLER et al. 2002). Diverging extension along the main Delamerian-Lachlan tectonic boundary eventually triggered the formation of the Tasman Fracture Zone after the first oceanic crust was formed between Australia and Antarctica in the Mid-/Late Eocene.

#### GONDWANA BREAKUP AND TRANSFORM/ PASSIVE MARGIN EVOLUTION

Prior to the rifting between Antarctic and Australia, the Ross Sea region and SE Australia were located in the hinterland of the Panthalassan margin of Gondwana (e.g., COLLINSON et al. 1994). Compressional tectonism across this continental margin was related to subduction and terrane accretion to the east of the present location of northern Victoria Land. Behind the Panthalassan margin developed a large basin system comprising the Transantarctic and Wilkes basins in Antarctica, and numerous basins that covered almost the whole Australian continent (e.g., VEEVERS 2006). Of the Australian basins the Eucla/Great Australian Bight, Duntroon, Otway, Bass, and Gippsland depositional centres (cf. MUTTER et al. 1985) may have been directly linked with the Antarctic Wilkes and Transantarctic basins (Fig. 1). Much of these massive sedimentary sequences are still preserved in Australia (cf. VEEVERS 2006 cum lit.) while remnants of late Paleozoic to Mesozoic deposition along the Pacific Antarctic margin are limited to the Beacon Group and confined to the vicinity of the Rennick Graben.

Nevertheless, this basin was once more extensive, and estimates of the minimum size of this basin extension as well as burial depths in northern Victoria Land rely on the correlation of the Kukri Surface with AFT data. A minimum time constraint for basin initiation can be derived from the extrusion of the ~350 Ma Gallipoli and Black Prince rhyolites. AFT ages from these superficial rocks and all outcrops east of the Rennick Graben are of Late Cretaceous to Paleogene time equivalent, and therefore indicate post-Carboniferous heating

of the Kukri Surface to temperatures >110 °C. Three of four zircon fission track ages from ~350 Ma Admiralty Granites from the boundary area between northern Victoria Land and Terra Nova Bay were reset to 220-250 Ma whereas seven of eight titanite fission track ages from samples of the same lithology, approximate altitude, and region give effective intrusion ages (FITZGERALD & GLEADOW 1988). These data suggest maximum paleotemperatures in the order of 200-250 °C for the Kukri Surface in northern Victoria Land. Substantially older AFT ages up to ~350 Ma of Kukri samples from the Eisenhower and Deep Freeze Ranges (BALESTRIERI et al. 1994, 1999) indicate that maximum paleotemperatures decreased towards the Terra Nova Bay region to <110 °C. When applying a Late Cretaceous/Paleogene paleogeothermal gradient of 25 °C km<sup>-1</sup> as assumed by FITZGERALD & GLEADOW (1988) and calculated by LISKER et al. (2006), such paleotemperatures refer to basin depths between 3 km (Terra Nova Bay) and up to 8 km (northern Victoria Land). Hence, the locus of the subsequent continental breakup underlies the region that links up to 15 km of Jurassic to Cretaceous sediments on the SE Australian shelf with up to 8 km of coeval overburden in northern Victoria Land. This approximate paleo-depocentre likely controlled the degree of crustal weakening and thus the focusing of extensional strain. Given a thickness of substantially less than 1000 m of Beacon strata below Ferrar rocks, only a minor section of this sedimentary column was deposited during the Permo-Triassic. Instead, the deposition rate must have increased subsequent to Ferrar emplacement, and maximum burial was likely reached near the present Antarctic / Australian margin in the Cretaceous, prior to the onset of Paleogene exhumation.

The AFT pattern changes towards west across the Rennick Graben where AFT ages of samples taken between Oates Land and Terre Adélie vary between ~100 and ~300 Ma (LISKER & OLESCH 2003, LISKER et al. 2006). This region was obviously not part of the Transantarctic Basin and its successor (George V Land, Terre Adélie), or buried substantially less (USARP Mountains/ Oates Land).

Neither of the two Pacific Antarctic margin sections divided by the Rennick Graben exhibits a distinct thermal signature related to the onset of Late Cretaceous Gondwana breakup. Instead, the timing of exhumation of northern Victoria Land, which was in the order of 4-8 km, coincides with the Eocene formation of the West Antarctic Rift System and the opening of the Tasman gateway (e.g., PFUHL & MCCAVE 2005). A later cooling/exhumation "event" is not recognized. In general, the long-term regional exhumation pattern suggest long-lasting E-W crustal extension and sediment deposition in a basin overlying both southeastern Australia and the northern Victoria Land region of Antarctica. A sudden increase of extension rates culminated in the Ferrar magmatic episode at ~180 Ma, and in the Cenozoic rifting of the Ross Sea. The latter was associated substantial faulting along parallel structures as the Rennick Graben. The Rennick Graben master faults and their continuation into Australia probably represent major lineaments that juxtapose crustal blocks of different rheological properties. In this context, Eocene exhumation results from uplift due to flexure and isostasy, followed and superimposed by thermal effects across the newly formed new margin (cf. LISKER 2002). We suggest initial margin formation between Antarctica and Australia as the result of shearing due

to the clockwise rotation of Gondwana, with different movement rates of both supercontinent fragments.

## CONCLUSIONS

AFT thermochronology represents the most important tool to unravel the exhumation history and long-term landscape evolution of northern Victoria Land and to conclude on Gondwana breakup and passive/transform margin evolution between Antarctica and Australia.

Paleo-isotherms derived from thermochronological data during the last two decades indicate the existence of a late Paleozoic-Mesozoic basin in northern Victoria Land and Australia. Intra-Gondwanan oblique extension and basin evolution lasted much longer than anticipated earlier, with sudden increase of extension being responsible for tectonomagmatic events within the basin, such as the ~180 Ma Ferrar event or Cenozoic West Antarctic rifting.

Thermal histories also reveal a characteristic pattern of increased exhumation and uplift since the Eocene for the region east of the regional lineament Rennick Graben – Tasman Fracture Zone – Woorndoo-Sorrel Fault Zone while substantially less exhumation occurred west of it. Increased exhumation associated with Early Cretaceous initial oblique rifting and Gondwana breakup between Antarctica and Australia is not observed. Instead, Eocene exhumation is likely linked with the rifting of the West Antarctic Rift System and/or the onset of sea floor spreading and the opening of the Tasman gateway.

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