

A STUDY OF GLACIAL-INTERGLACIAL VARIATIONS OF THE MARINE STABLE CARBON ISOTOPE RECORD USING A NON-REDFIELD BIOGEOCHEMICAL MODEL



PAL MOD

GERMAN CLIMATE MODELING INITIATIVE



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Motivation

Photosynthetic plankton has a lower ^{13}C concentration than the aqueous CO_2 from which it was formed. The isotopic discrimination during photosynthesis increases the $^{13}\text{C}:^{12}\text{C}$ ratio (or $\delta^{13}\text{C}$) of dissolved inorganic carbon (DIC). Here, we investigate the effect of two different parametrizations of biogenic fractionation on the carbon-isotopic signature of DIC under present and glacial climate conditions.

Biogenic carbon fractionation

We consider two parametrizations of biogenic fractionation. Rau (1994) found that the isotopic depletion of phytoplankton $\delta^{13}\text{C}_p$ increases with the availability of aqueous carbon dioxide CO_2^* :

$$\delta^{13}\text{C}_p = -a_0\text{CO}_2^* - b_0$$

Laws et al. (1997) found that $\delta^{13}\text{C}_p$ also depends on the isotopic composition of CO_2^* (depending on temperature and CO_3^{2-}) and photosynthesis μ :

$$\delta^{13}\text{C}_p = \delta^{13}\text{C}\text{CO}_2^* - (a_1\text{CO}_2^* + b_1\mu) / (a_2\text{CO}_2^* + b_2\mu)$$

Conclusion / Summary / Outlook

Different parametrizations of biogenic fractionation lead to discernable changes in the carbon-isotopic composition of DIC. The differences are seen in the entire water column. This may be particularly an issue in future model - data comparisons for the glacial ocean.

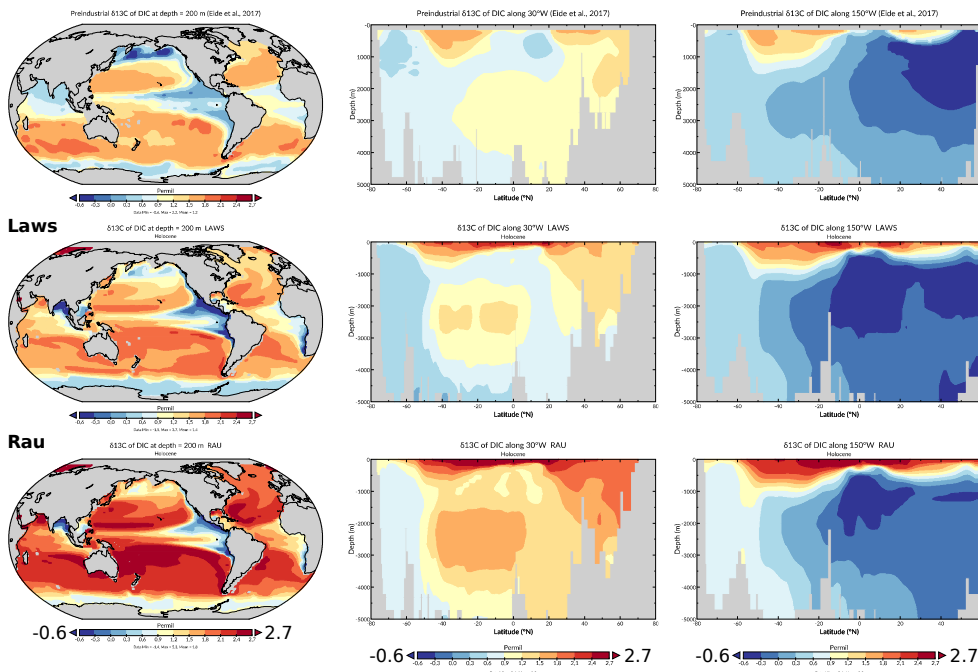
Holocene

Reconstruction

200 m

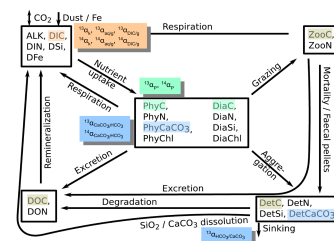
Atlantic

Pacific



Model setup

We use the Regulated Ecosystem Model RECOM which does not rely on fixed Redfield ratios for organic soft tissue. Instead, the C:N and C:Chl ratios in phytoplankton respond to light, temperature and nutrient supply. We consider isotopic fractionation during uptake and dissolution of CO_2 , calcification, and photosynthesis of phytoplankton, plus radioactive decay of ^{14}C .

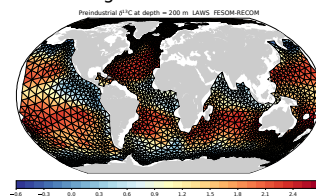


Tracers are transported via the MITgcm, forced with climatological fields derived in fully coupled climate simulations for the Holocene (Wei & Lohmann, 2012) and the Last Glacial Maximum (Zhang et al., 2013). Our MITgcm-RECOM simulations build upon previous work by Völker & Köhler (2013). Dust fields are by Albani et al. (2016). Model resolution is 2° longitude x 0.38° to 2.0° latitude x 30 levels. Integration time (so far): 2 - 3 kyr.

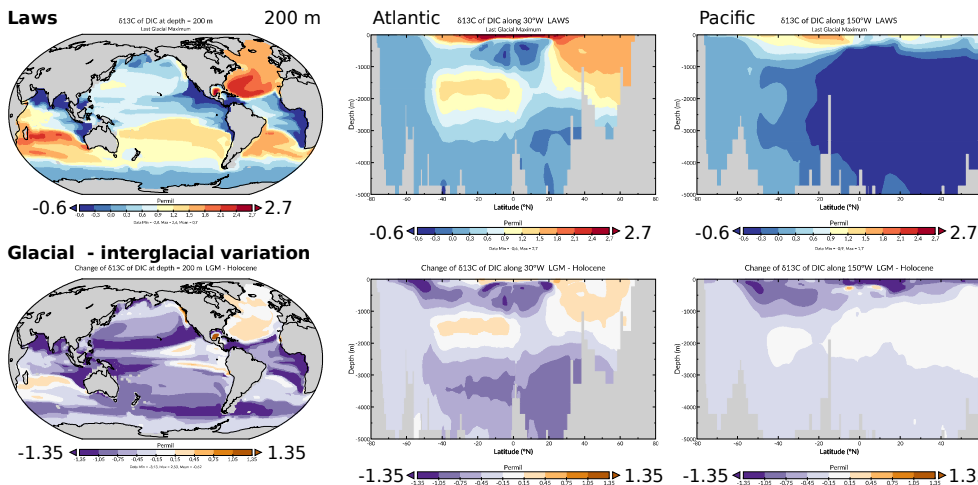
The biogenic fractionation parametrization according to Rau leads to higher values of $\delta^{13}\text{C}$ of DIC than the parametrization by Laws et al. The elevation is not limited to the euphotic zone but is also obvious in deeper waters. In our Holocene simulations the parametrization according to Laws et al. leads to better agreement with the reconstruction by Eide et al.

Current model development

RECOM has become part of the most recent version of the AWI climate model. The AWI climate model features FESOM, a multi-resolution sea ice-ocean model solving the equations of motion on unstructured meshes. The figure below gives an early impression of a test run with prescribed preindustrial climate forcing.



Last Glacial Maximum



References

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We find isotopic depletion of DIC except for small areas such as the subtropical North Atlantic and the Indian.

