## Augmenting the BSH operational forecasting system by in situ data assimilation

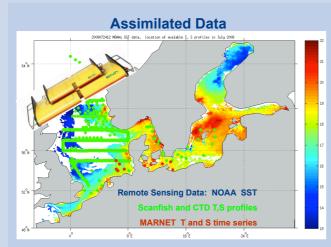
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## **Abstract**

Quality of the forecast provided by the German Maritime and Hydrographic Agency (BSH) for the North and Baltic Seas had been previously improved by assimilating satellite sea surface temperature SST (project DeMarine, Losa et al., 2012). We investigate possible further improvements using in situ observational temperature and salinity data: Marnet time series and CTD and ScanFish measurements. To assimilate the data, we implement the Singular Evolutive Interpolated Kalman (SEIK) filter (Pham et al., 1998). The SIEK analysis is performed locally (Nerger et al. 2006) accounting for/assimilating the data within a certain radius. In order to determine suitable localisation conditions for Marnet data assimilation, the BSHcmod error statistics have been analysed based on LSEIK filtering every 12 hours over a one year period (September 2007 – October 2008) given a 12-hourly composites of NOAA's SST and under the experiment conditions corresponding the maximum entropy. The principle of Maximum Entropy is also used as an additional criterion of plausibility of the augmented system performance.



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The data archive is based on measurements collected by BSH, Sweden's Meteorological and Hydrological Institut (SMHI) and the Institute of Marine Research (IMR, Norway)

## **Principle of Maximum Entropy**

general formulation, Kivman et al., 2001

 $He(\rho) = -\int_{x} \rho(x|y) \ln \frac{\rho(x|y)}{\mu(x)} \prod dx$ alisations x given the y or the conditional PE  $\rho(x|y)$  is the probability density function (PDF) of model trajectories al PDF also called the analysis PDF, which expresses the state of our know knowledge about the model state when data are observed, u(x) is the lowest information about x. The maximum probable x or mean with respect to  $\rho(x/y)$  is  $x_i = M_m x_m + M_d x_d$ 

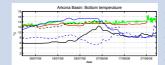
 $M_{...} = L_{\bullet}L, \qquad M_{d} = H_{\bullet}H$ M are nonnegative, M reflect our assumptions on the model and data error covariances. Operators  $M_m$  and  $M_d$  are nonnegative,

self-adjoint and  $M_m + M_d = I$ . M is an operator-valued measure. Given  $\lambda_i$  of  $M_m$  or  $M_d$  matrixes, one can calculate  $He(M) = -trace(M_d \ln M_d + M_m \ln M_m) = -\sum_{i=1}^{N} [\lambda_i \ln \lambda_i + (1 - \lambda_i) \ln(1 - \lambda_i)]$ 

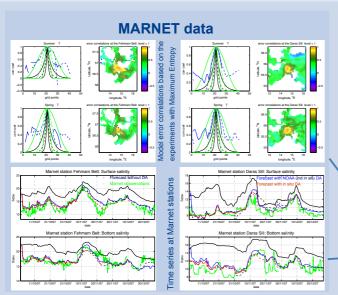
In Kalman type Filtering

The maximum probable x or state vector analysis  $x^a$  is  $x(t_n)^a = x(t_n)^{f,m} + K_n(d_n - Hx(t_n)^{f,m})$ , where  $x(t_n)^a$  and  $x(t_n)^f$  denote analysis and forecast of the model state at certain time  $t_n$ ,  $y_n$  is observations available at  $t_n$ , K is the Kalman gain  $K_{-} = P_{-}^{f}H(HP_{-}^{f}H^{T} + R)^{-1}$ 

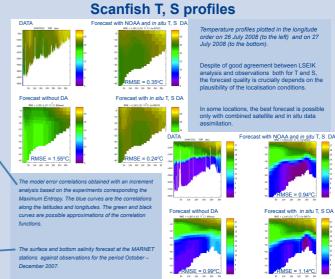
Here, following Pham (1998),  $P_n^f$  is the forecast error covariance matrix, H is the observation operator and R is the observational error covariance matrix. To calculate the entropy  $He(\rho)$ , we just need to know  $\lambda_s$  of the Kalman gain matrix (using SVD decomposition). Such a matrix could be constructed by collecting and considering  $K_s$ , for instance, globally over a certain period of time or locally. The last variant is valuable for validation of localisation conditions.



Temporal evolution the bottom temperature forecast at the MARNET station "Arkona Basin" produced with BSHcmod without DA (black); with LSEIK analysis of the model and NOAA's SST DA under statistical conditions corresponding the Ite=4.86 for the period 25 June – 8 August 2008 (blue solid); based on NOAA's SST LSEIK analysis underror statistics with It=271 for the same period (blue dashed); assimilating satellite SST and in situ T, S data including MARNET (black dashed); assimilating only in situ fulls data (red.). The oreen curve depicts



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