

Bachelorarbeit

Zur Erlangung des akademischen Grades Bachelor of Science

Thema der Arbeit:

Thermokarst lake dynamics on the Alaska North Slope and their influence on
organic matter deposition

eingereicht von: Linda Siegert
Matrikelnummer: 582443
E-Mail: siegertl@hu-berlin.de

Gutachter/innen: Prof. Dr. Christoph Schneider
Dr. Josefine Lenz

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Bachelor's thesis

To earn the academic degree Bachelor of Science

Topic of the thesis:

Thermokarst lake dynamics on the Alaska North Slope and their influence on organic matter deposition



Photo: J.Lenz

Submitted by: Linda Siegert
Matrikelnummer: 582443
E-Mail: siegertl@hu-berlin.de

Internal Supervisor: Prof. Dr. Christoph Schneider
External Supervisor: Dr. Josefine Lenz

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Abstract

Arctic regions are affected by the warming twice times faster than the global mean. (Overland et al., 2017) Within the last years permafrost has rapidly thawed which is a significant issue in relation to climate change and landscape structures in arctic regions.

The biggest issue of rapidly thawing permafrost are the stored carbon sources in the ground which contain a lot of methane and carbon dioxide. Thus climate change is a lot influenced by self-enforcing feedback effects which have a very high impact (Pithan et al., 2014).

This study is based on 10 short sediment cores taken from thermokarst lakes on the Alaska North Slope in July and August 2018 by researchers of the Alfred Wegener Institute for Polar and Marine Research and the University of Alaska Fairbanks.

The main focus of this thesis is the understanding of the dynamics of thermokarst lakes based on geochemical analyses, physical parameters and remote sensing imageries.

Within this study the dynamics of thermokarst lakes were discussed and researched in relation to their decomposed organic matter content.

The relation between the lake size, water depth, lake ice type and organic matter input was discussed as well.

Lastly this thesis canvased the influence of salinity of organic matter input.

The method of these classifications is qualitative and is handled by data tables without statistical classifiers.

1. Introduction

The northern part of Alaska, namely the Alaskan North Slope, is characterized by arctic conditions in all categories. It is especially influenced by continuous permafrost. Permafrost is vulnerable to climate change and as a consequence of increasing air temperatures thawing processes create visible changes in the landscape because of the resulting increased ground temperatures as well (Romanovsky et al. 2010).

An obvious evidence for thawing processes and the resulting degradation of permafrost are thermokarst lakes.

These types of lakes are very complex in their development and very relevant on a global level since they act as carbon sinks or carbon sources and even provide information about the global greenhouse gas emissions according to their state of development. (Walter et al. 2007)

The velocity of thawing is very significant according to the release of greenhouse gases. Abrupt thawing causes more degassing of carbon dioxide (CO₂) and especially methane (CH₄) than gradual thawing over years (**Figure 1**).

It is also necessary to realize that the lakes start to freeze every autumn and thaw every spring. In relation to the cycle of the seasons the albedo changes as well. If the thawing processes are advantaged more through increasing temperatures, the annual albedo is higher which means a self-enforcing feedback effect. So it is very important to study the dynamics of the lakes to provide information about the impacts of the climate change in relation to permafrost thawing processes. Deposited sediments of the lake ground can also give a view on the variability of the environment during the Holocene. (Lenz et al. 2016)

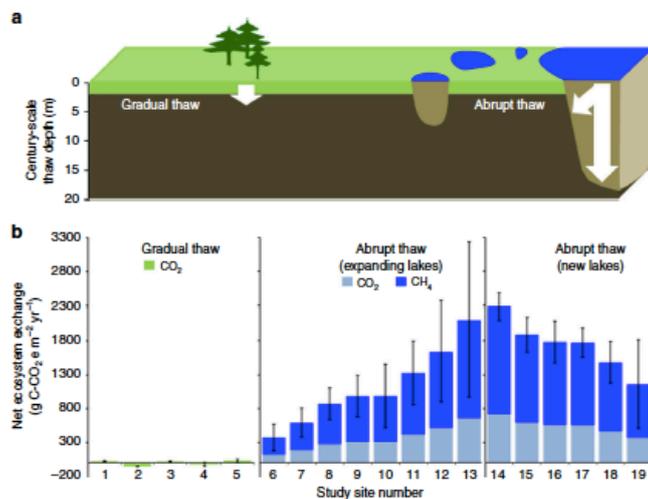


Figure 1: Difference of gradual and abrupt thawing. **a** schematic, **b** Net ecosystem exchange in Greenland (1,2), Healy, Alaska (3-5), western Alaska (6-9), interior Alaska (10, 11, 14-19), northeast Siberia (12, 13), (Walter Anthony et al. 2018)

1.1 Thermokarst lake development

The periglacial landscape is characterized by ice-rich permafrost and ice wedges.

Thermokarst lakes develop by ground ice loss, ground subsidence and pooling of water. That is why the natural rhythm of summer and winter is important by the development of thermokarst lakes, too. The active layer of the permafrost starts to thaw at the beginning of summer and starts to freeze at the end of the late summer.

The abrupt thawing causes visible bulges in the landscape (**Figure 2 (b)**). The thawing causes a depression which gets filled by thawing waters and precipitation waters (**Figure 2(c)**). Ice wedges accelerate these processes because of the ice-richness which indicates more thawing waters. (**Figure 2 (d)**). Ice wedges also cause degrading in a much more intense way at the places where they developed. As ice wedges develop in a high amount in arctic tundra, the occurrence of them is very significant. The non-frozen Talik also causes thawing around the lake because of the heat flux in direction to the surrounding permafrost (**Figure 2 (e)**), Grosse et al., 2013).

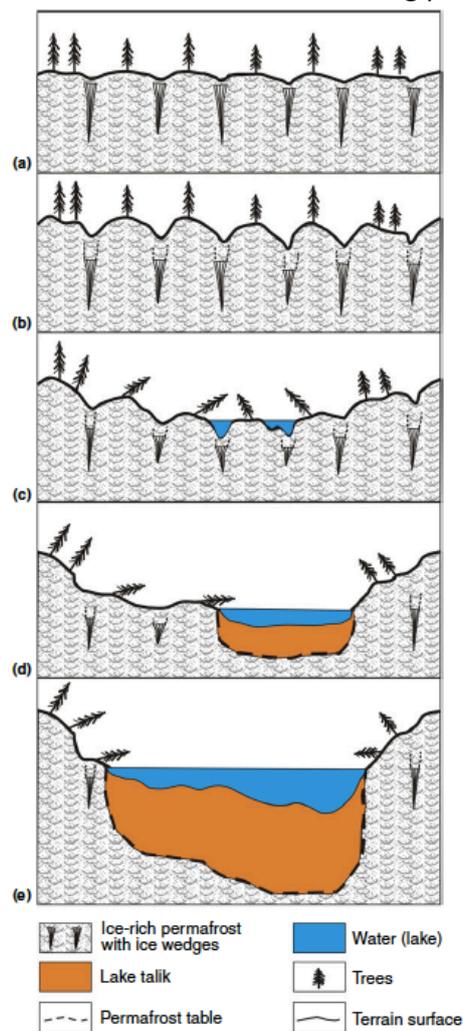


Figure 2: Thermokarst lake development in five stages, (Grosse et al., 2013)

1.2 Research hypotheses

The main focus of this bachelor's thesis project will be the analyses of the relationship between the lake type characteristics and organic matter deposition of lake sediments based on 10 short sediment cores with references to their dynamics.

One category of the lake classification will be dynamic and stable lakes. For this category comparison of remote sensing imageries will be used to get an idea of the lake size changes over the past 32 years. It is also expected that the stable lakes consist of more decomposed organic matter than the more dynamic lakes because they tend to change in size more often which is a reason for a higher rate of terrestrial plants than aquatic plants. It is expected as well that larger and deeper grounded ice lakes contain less organic matter input and rather feature aquatic plants than smaller and shallower floating ice lakes.

Marine influenced and freshwater lakes will be studied to get an idea of the organic matter in relation to their salinity.

Consequently, the three main research hypotheses are:

1. Sediments of stable lakes comprise more organic matter decomposed and rather aquatic plants than dynamic lakes which may contain less organic matter decomposed and rather terrestrial plants.
2. Sediments of smaller, shallower, grounded ice lakes have a lower organic matter input than larger, deeper, floating ice lakes.
3. Lakes with higher salinity provide a lower organic content than lakes with lower salinity.

2. Study area

The study area extends from 70.4° to 70.9° N latitude and from -152.2 to -154.5° W longitude (**Figure 3**).

The region is in the North of Alaska on the Alaska North Slope and is also partly located on the Outer Arctic Coastal Plain (OACP). Thus, this area is influenced by marine waters. Especially the north coast of Alaska is influenced by a very fast coastal erosion which influences the salinity and drainage of thermokarst lakes as well. In a span of 1955-2005 the coast has eroded up to 0.9 km at some points in the north of Teshekpuk lake (Mars et al, 2007).

22.5 % of the OACP are covered up with thermokarst lakes and 61.8 % of the landscape are drained thermokarst lake basins. (Jones et al., 2015)

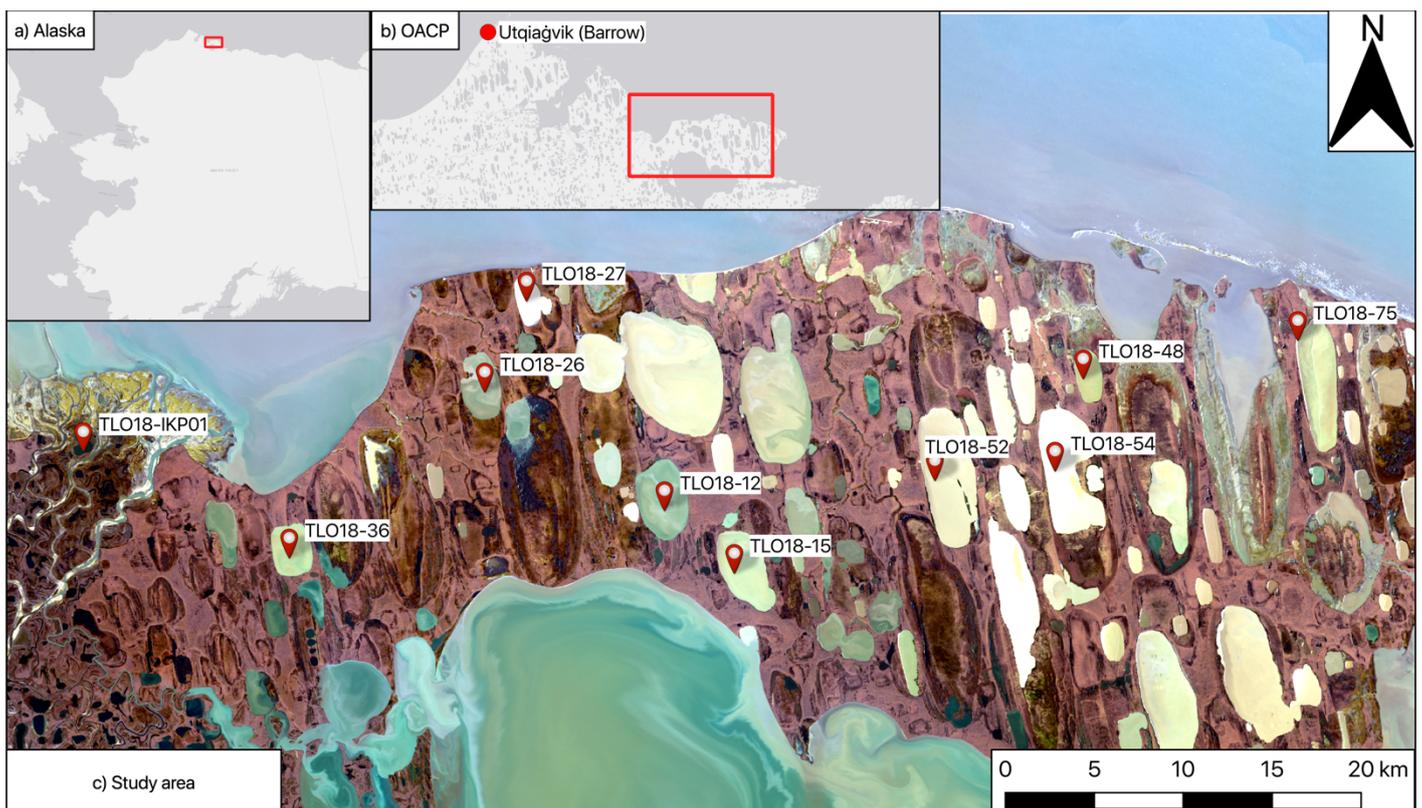


Figure 3: Spatial classifying of the study area, **a)** Study area (red) marked at the North of Alaska, **b)** Study area (red) marked on OACP, **c)** Concrete study area with taken cores, (Backmap a) and b) from Esri (Copyright © Esri), c) Landsat 8 satellite imagery provided by USGS Data Center)

Ten cores were taken on 29th and 30th of July in 2018. Every core was taken from a different lake. The core length was measured right after coring.

Water depths and water temperatures were measured as well. The water depths reach up from 67 cm to 350 cm which shows a variability of lake depths.

Water temperatures are very similar from 13 °C to 16.5 °C.

Surface water pH values and bottom water pH values were measured to analyse the water quality of the lakes. They reach from a minimum of 6.99 to a maximum of 8.17. (**Table 1**)

Pure water has a pH value of 7 which means that the lakes have no acidic properties. They rather tend to be more basic which means the lake waters contain more free hydroxyl ions.

Basic pH values in lake waters can be found because of a higher biotic productivity. (Institut Dr. Flad)

Table 1: Core and lake conditions during field work

Core ID	Latitude (°N)	Longitude (°W)	Date of collection (DD.MM.YYYY)	Core length after coring [cm]	Water depth [cm]	Water T [°C]	Surface water pH	Bottom water pH
TLO18 ¹⁾ -12	70.76377	153.56076	29.07.2018	40	250	13	8.10	7.77
TLO18-15	70.73219	153.45337	30.07.2018	42	220	14	7.92	7.62
TLO18-26	70.82292	153.83806	29.07.2018	62	230	16	6.99	6.99
TLO18-27	70.86937	153.77577	29.07.2018	14	64	16	7.84	7.19
TLO18-36	70.73708	154.13290	30.07.2018	35	200	13.5	8.02	7.62
TLO18-48	70.83154	152.92044	30.07.2018	31-33	160	14.0	8.17	7.54
TLO18-52	70.77983	153.14758	30.07.2018	22	200	-	7.63	7.43
TLO18-54	70.78470	152.96373	30.07.2018	13	100	-	7.49	-
TLO18-75	70.85054	152.59093	30.07.2018	15.5	100	16.5	7.80	7.46
TLO18-IKP01	70.78878	154.45096	29.07.2018	58	350	15	8.03	7.81

¹⁾ TLO18 denote Teshekpuk Lake Observatory 2018

2.1 Permafrost

Permafrost is classified as a

The landscape is marked by the influence of the continuous ice-rich permafrost up to thicknesses of minimum 410 m (Jorgenson et al. 2008).

The temperatures of the permafrost are cold with mean annual values of -6 to -10 °C (**Figure 4**, Romanovsky et al. 2010). Hence this region consists of a lot of thermokarst lakes because of the thawing processes caused by summertime temperatures and because of the intensified feedback effect by the climate change (Pithan et al., 2014).

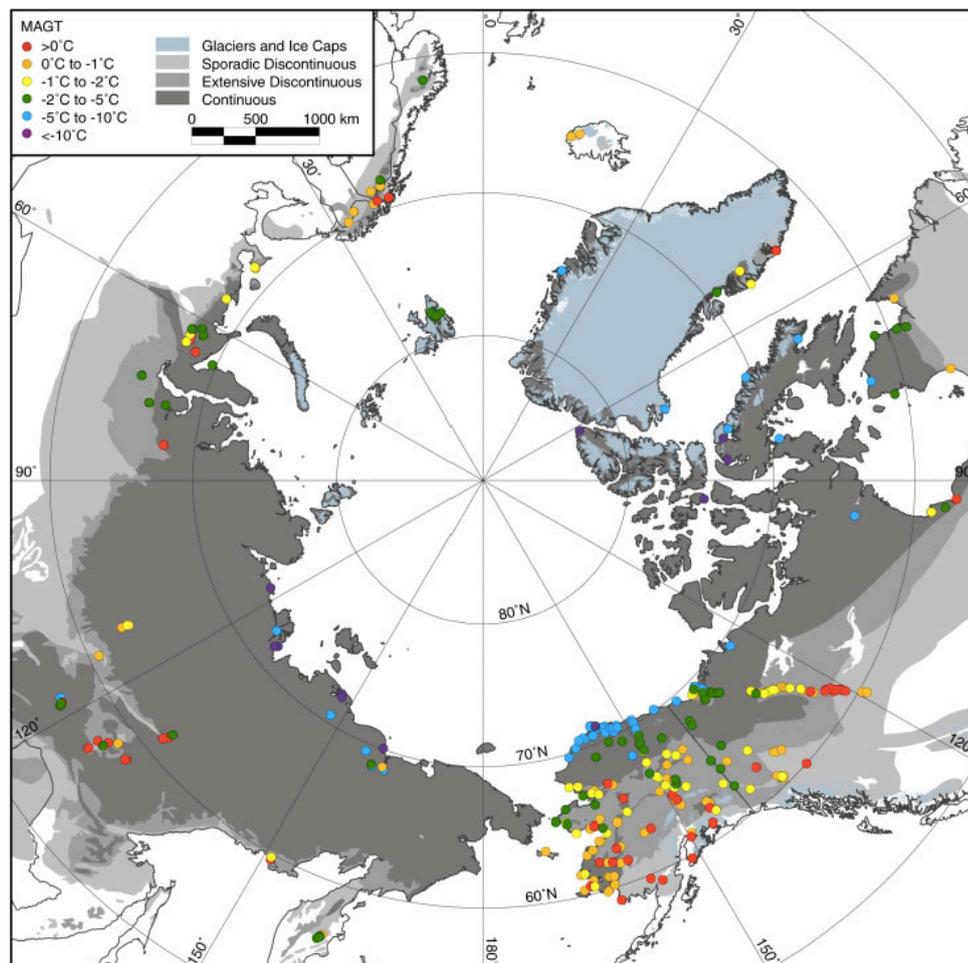
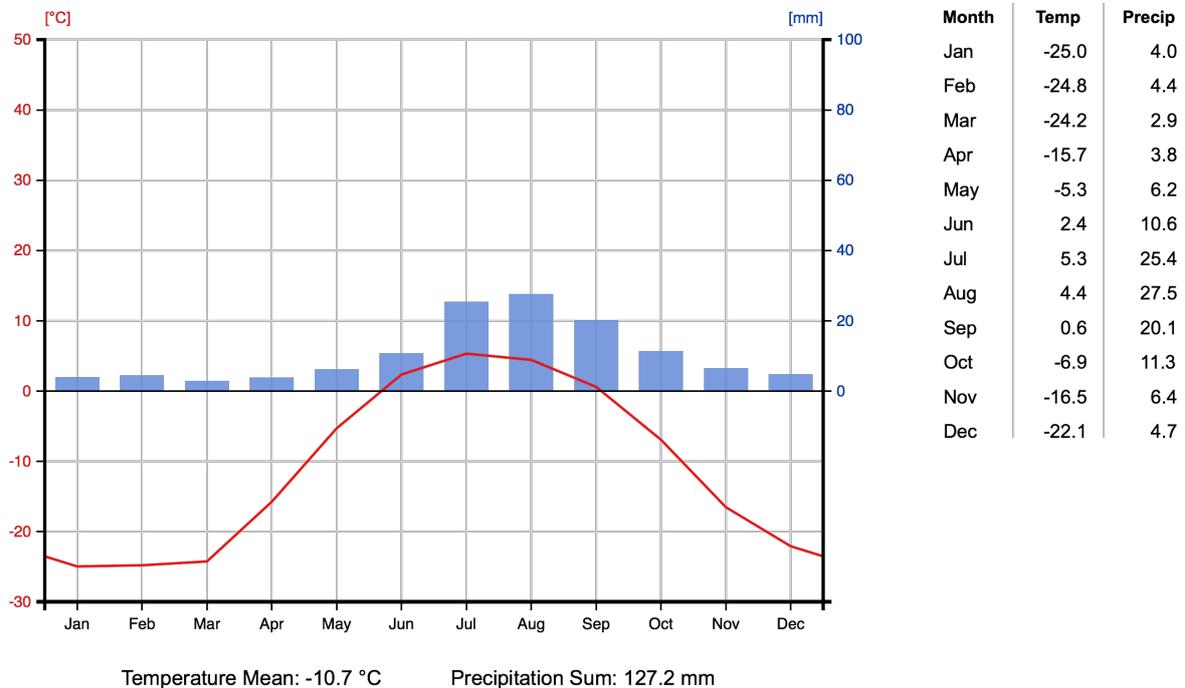


Figure 4: Permafrost classifications and permafrost temperatures (Romanovsky et al. 2010)

2.3 Climate and vegetation

Barrow/W. Pos, United States Of America

71.3N, 156.78W | Elevation: 4 m | Climate Class: ET | Years: 1987-2016



Data Source: www.ncdc.noaa.gov/ghcnm/

© ClimateCharts.net

Figure 6: Climate chart of Barrow, Alaska from 1987 to 2016 (NOAA)

The study area is influenced by the arctic climate which is why there are annual mean air temperatures of -10.7 °C. The summer months are cool with a mean July air temperature of 5.3 °C in Barrow and winter months can get very cold with mean January air temperatures of -25 °C. There are very low precipitation rates in this study area of about 127.2 mm/a in average which indicates semi-arid conditions of the study area (**Figure 6**, Shulski and Wendler 2007).

The vegetation is mainly wetland vegetation with wet graminoids and moss. Because of the permafrost influence the landscape is dominated by high- and low-centered polygonal tundra (Nowacki et al. 2002). Especially because of the high amount of lakes in this area there is a high occurrence of lacustrine algae and marine algae in littoral parts as well.

3.Methods

3.1 Field work

The field work was conducted during an expedition of Alfred Wegener Institute for Polar and Marine Research (AWI) and also of the University of Fairbanks Alaska (UAF). It took place in July and August of 2018. This expedition was a part of the PETA-CARB project and was led by Prof. Dr. Guido Grosse (AWI, University Potsdam) and Dr. Benjamin Jones (UAF) based out of Teshekpuk Lake Observatory.

The main focus of the expedition was the submarine permafrost, the vegetation types in arctic tundra regions and also dynamics of thermokarst lakes.

Dr. Josefine Lenz (AWI, UAF) drilled 10 cores from thermokarst lakes using a piston corer while standing on a float boat. (**Figure 7**) The lengths of the cores reach from 12 cm to 61.5 cm. It was cored until the frozen ground was reached which is why there was no ice contained in the cores. Every core was drilled from a different lake in order to assess spatial variability. The cores were almost always taken from the central of the lakes.

Lake depth, surface water temperature and electric conductivity were measured by Dr. Josefine Lenz as well.

The cores were transported to AWI and were stored by 4 °C until November 2018 when they got opened.



Figure 7: Core taking on a float boat with a piston corer (Photo: Juliane Wolter)

3.2 Laboratory work

To ascertain the characteristics of the deposits of the lake ground in relation to carbon content, sediment compounds and vegetation classes the samples got analysed on a geochemical base. The cores were subsampled to determine the total carbon (TC), total organic carbon (TOC), total nitrogen (TN) and isotope ratio of total organic carbon ($\delta^{13}C$). Besides the geochemical analyses the cores were described by their visual characteristics (amount of visible organic), their colour, their grain size and their size.

3.2.1 Sample preparation

The cores were cut in halves with an electronic saw. One half was used to archive each core and was stored isolated. The other half was used to get subsamples taken from it. The samples got taken with a syringe. This tool helped against a contamination of the samples under each other. The taken subsample was split in two 12.5 ml plastic jars (A and B sample) thus every sample was archived as well and the other part could be used for the geochemical analyses.

The studied core half was described visually before sample extraction. Every core was classified by its length, colour, grain size, organic matter content and structure. The colour was defined with the Munsell colour system by Albert H. Munsell. The grain size was conducted qualitatively with roll validations. The organic matter content was classified visually, as well as the structure of the cores.

The wet samples got weighted and were freeze-dried afterwards to extract the water.

The dried samples got milled with a Fritsch pulverisette 5 planetary mill to create a homogenous powdered sample. Each sample was milled in agate jars and agate marbles for 5 minutes at 360 rotations per minute and was transferred into plastic jars with a spatula and a brush.

For this study it was relevant to use the sample data of the first 10 cm averaged of each core. This decision was made because there is no relation to a specific point of time since there was no radiocarbon dating measurement made. This is why there is a discussion about the lake classification at the end besides all other named work hypotheses.

3.2.2 Total carbon and total nitrogen

Every sample was measured twice for the analyses of TC and TN, which both could be determinate within one measurement. The amount of each sample weighted between 5 to 5.8 mg. The samples were weighted with a Sartorius micro M2P laboratory scale which has an accuracy of ± 0.001 mg. The samples were put in tin capsules and also Tungsten(VI)Oxide was added to catalyse the combustion process.

The vario EL III Elementar Analyzer (Elementar Analysensysteme GmbH) accomplished the measurement process. It heats the samples up to 950 °C with a helium atmosphere which is oxygen-saturated. At 850 °C the CO₂, N₂, NO and NO₂ are formed and are reduced in a reduction tube.

The actual measurement begins at 50 °C which causes a mobilisation of nitrogen and a transfer to the measuring chamber. The nitrogen is getting detected by a heat conductivity sensor which appears as a peak in heat conductivity.

Afterwards, the vario EL III heats up to 130 °C again to begin the carbon dioxide transfer to the measuring chamber as well which causes another peak of heat conductivity. To prove the extent of the peaks it was important to put these into context with the sample weight. Resulting values of TC and TN were given in wt% with an accuracy of 99.95 % for carbon and 99.9% for nitrogen.

It was necessary to prove these accuracies with calibration standards. Acetanilide, sucrose and 30 % EDTA were measured at the beginning of the set to control the set of standards used in between the sample measurements after every 30 measurements consisting of 20 % EDTA, 12 % calcium carbonate and two soil standards. Even the standards were measured twice and could have only a measurement difference maximum of 10 %.

3.2.3 Total organic carbon

It is important to distinguish total organic carbon from the total carbon content of the samples because the total carbon content also contains the inorganic carbon like little shell remains. The total organic carbon was measured with the elementary analysis device varioMAX C Element Analyzer (Elementar Analysensysteme GmbH). It functions with catalytic tube pyrolysis while pure Nitrogen (99.996 %) is the carrier gas. The amount of the measured sample is based on its total carbon content. The samples were weighted with a Mettler Toledo XS105 dual range analysis scale with an accuracy of ± 0.1 mg) in steel crucibles.

At the beginning of the analysis, a set of calibration standards composed of pure glutamate, 30 % glutamate and 2:3 glutamate were used. After every 15 samples 2:3 glutamate, 10:40 glutamate, 5:45 glutamate and 1:19 glutamate were used to control the measurements.

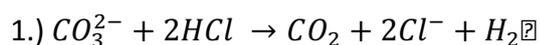
The samples were heated up to 580 °C in a helium atmosphere which is oxygen-saturated to form them in carbon dioxide formation. To assure the complete carbon dioxide formation the gases were heated up to 930 °C. The device exhibits a detectable carbon content minimum of 0.1 %. The carbon dioxide peak during integration is determined by the cut off value which is calculated by sample mass and anticipated carbon amount. The total organic carbon content is defined by the value and arising of the peak. The accuracy of this method amounts to 99.9 %. Within the values of TOC and N there was calculated a ratio which implies the state of decomposition of the organic matter in the samples. The organic matter qualifies a better conserved state, if the values are higher.

3.2.4 Stable carbon isotopes

Carbon ions needed to be removed for stable isotope analysis. For this, a little amount of each sample was filled in Erlenmeyer glass flasks. Every sample was enriched with 20 ml 1.3 molar hydrochloride acid (HCl). The solution was heated up to 97.7 °C on a hot plate for three hours. They were filled up with purified water afterwards. The samples were decanted every day for three days in a row to remove the Cl⁻ ions. It was necessary to reduce the Cl⁻ content of under 500 ppm. The elimination of these ions was needed to avoid disorders in the electromagnetic field of the mass spectrometry which is used for the measurements. Divergent ionization reacts could be avoid with this process as well. The samples were tested with Quatofix Chloride test stripes to ensure the amount of Cl⁻ ions in every sample.

The samples were vacuum-filtered, dried at 50 °C and hand-grinded afterwards to get the samples decarbonised and pulverised again.

The following equation shows the reaction while decarbonising the samples:



To prepare the samples for the measurement it was necessary to weigh in a targeted amount of the samples with the following equation:

$$2.) \text{target weight [mg]} = \frac{20}{[\text{wt}\%]}$$

The weighing was conducted with a Sartorius micro M2P scale (accuracy of ± 0.001 mg) with a maximum deviation of 0.05 mg from the target weights.

The analysis was accomplished by Delta V Advantage Isotope Ratio MS supplement (Thermo Fisher Scientific) with a Flash 2000 Organic Elemental Analyzer (Thermo Fisher Scientific). Both devices use helium as carrier gas.

The first step of the analysis is the sample combustion and oxidation at 1020 °C with chromium dioxide as an oxidant which causes the formation of carbon dioxide (CO₂) and pure nitrogen formations.

Elemental copper causes a reduction to pure nitrogen (N₂) at 650 °C and leaves the CO₂ unmodified. The gas chromatography tube separates the gases. Nitrogen comes into the mass spectrometry faster because it is lighter than CO₂. Inside the tubes, external nitrogen and external CO₂ pass through the mass spectrometry to measure reference values. The measuring in general works by ionising the samples within electron impulses which are caused by energy inducing. The ions get separated by detection in the analysis unit where they get categorised by their mass and charge ratio and their energy intensity. The raw measured data has been calibrated to standards using a linear correction. The 1 σ standard error is better than ± 0.15 ‰.

3.3 GIS analytics

3.3.1 Lake size

For lake size dynamics it was necessary to compare older stadiums with the latest situation of the lakes. Therefore, remote sensing images were used to compare the lakes size visually. To find suitable images Google Earth Engine was utilised to find satellite images due to the summer month. The cloud cover was filtered, too. As a result of this method, a Landsat 5 image from 12th August 1986 and a Landsat 8 image of 19th September 2018 were used to digitalise the lake outlines. The digitalisation was produced with QGIS3.6 and its polygon tool. Afterwards the lake size could have been ascertained with field calculator. For Lake IKPO1 there was only one suitable image from 25th July 2006 available to show the dynamics of the lake properly.

3.3.2 Grounded ice lakes and floating ice lakes

The categorisation of grounded ice lakes and floating ice lakes was made with the assistance of scientific papers by Dr. Christopher D. Arp et al. from 2011 and 2012, hence these studies show already ascertained classifications for this study area.

4. Results

4.1 Field work

First results were made while field work. Important results of field work were the core taking and the measurement of electrical conductivity of the surface water.

The measured electrical conductivity of the surface water had a wide range of values between 143 $\mu\text{S}/\text{cm}$ and 10380 $\mu\text{S}/\text{cm}$. (Fig.) The lakes 48, 75 and IKPO1 have the highest values with values between 7000 and 10380 $\mu\text{S}/\text{cm}$. The lowest values with 143 to 168 $\mu\text{S}/\text{cm}$ were measured in the lakes 54 and 52. The remaining lakes have very similar values from 400 to 570 $\mu\text{S}/\text{cm}$.

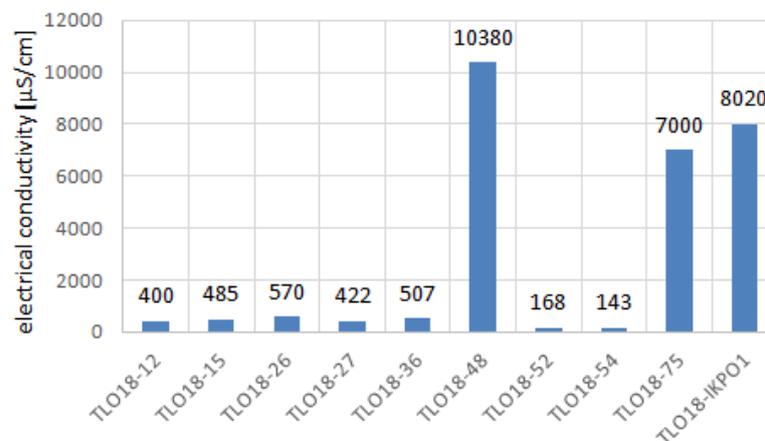


Figure 8: Electrical Conductivity of lakes

4.2 Laboratory work

4.2.1 Core descriptions

The cores were described by their length, colour, grain size, organic matter content and structure (Table). The cores TLO18-27, -52, -54, -75 are the shorter cores of this study, they have a length of 12 cm to 22 cm. The longer cores TLO18-12, -15, -26, 36, -48 and -IKPO1 reach up from 32 cm to 61.5 cm. The cores spread from yellowish tones to reddish and brownish tones and also from lighter to black tones. The grain sizes reach from clay to sandy material. Every core contains silt from fine to coarse. The cores TLO18-15, -36 and -48 contain coarser grain sizes up to fine sand and coarse silt. The other cores are more clayish and contained fine to middle silt.

TLO18-26, -36, -48 and -IKPO1 are organic dominated with visible organic remains, the others are minerogenic dominated. The organic dominated cores show different rates of decomposition. TLO18-26 and -36 exhibit well decomposed organic, while TLO18-48 contains macro organics and TLO18-IKPO1 even exhibits less to moderately decomposed organics.

Table 2: Core descriptions

Core ID	Length [cm]	Colour [Munsell, 1994]	Grain size	Organic matter content	Structure
TLO18-12	40	2.5Y 4/2, -> 2.5Y 3/1	fU to T	fine organics at the end, lighter layers contain coarser organic	layering visible, minerogenic dominated
TLO18-15	42	2.5Y 4/3 -> 5Y 3/2 -> 5Y 3/1	fS -> gU->T	visible organics at the bottom	minerogenic dominated
TLO18-26	61.5	5y 2.5/2	fU	well decomposed, coarser remains	organic dominated, well layered
TLO18-27	13	5y 4/2	mU -> T/fU	no organic visible	very minerogenic dominated, very dense
TLO18-36	36	5y 3/2 -> 5y 3/1	fU -> gU	well decomposed	organic dominated, coarser minerogenic compounds at the bottom
TLO18-48	32	5Y 3/2 -> black -> 5Y 3/1 -> 5Y 4/1	T/fU -> fU -> gU -> T	no organic visible at the top, macro organics visible in the middle	organic dominated, fine layered, minerogenic at the bottom
TLO18-52	22	5Y 3/1	mU	no organic visible	very minerogenic dominated (silty), clay laminae
TLO18-54	12	2.5Y 4/2 -> 5Y 4/1	fU -> T	no organic visible	very minerogenic dominated (clay), very dense
TLO18-75	16	2.5Y 2.5/1 -> 2.5Y 4/1	T to fU	poorly decomposed living vegetation at the top	minerogenic dominated, very dense
TLO18-IKPO1	57	2.5Y 3/3 -> black -> 2.5Y 3/2 -> 2.5Y 2.5/1 -> 2.5Y 3/2 -> black	mU	larger remains, possibly aquatic, less-moderately decomposed	organic dominated

*-> = colour and grain size changes during core

4.2.2 Total carbon and total nitrogen

TC and TN values were measured in one measurement. The TN values are significantly lower than the measured TC values. The values of TN and TC correspond with each other which shows a relation between total carbon and total nitrogen content. The TN values range between 0.21 wt% to 0.674 wt%. TLO18-12, -26 and 36 have the highest values with 0.56 wt% to 0.674 wt%. The lowest values have been measured in TLO18-27, -54 and -75 with 0.27wt%, 0.281 wt% and 0.21 wt% (Figure).

The TC content is significantly higher and reach from 3.89 wt% to 11.487 wt%. The highest values were measured in TLO18-12, -26, -36 and -48. These values range from 7.04 wt% to 11.487 wt%. The lowest were measured in TLO18-27, -52, -54 and -75 with 4.35 wt%, 4.82 wt% , 4.5705 wt% and 3.89 wt% which is barely a third of the maximum (Figure).

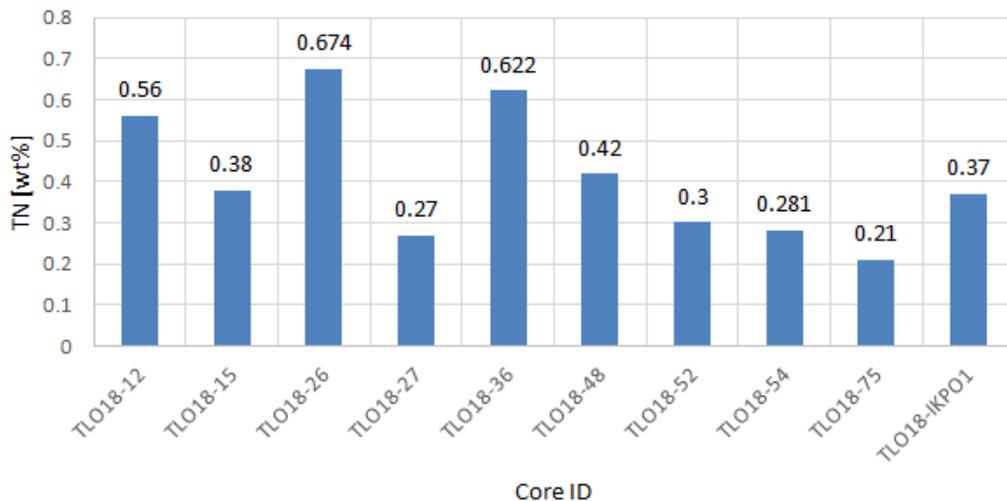


Figure 9: Total nitrogen of all cores

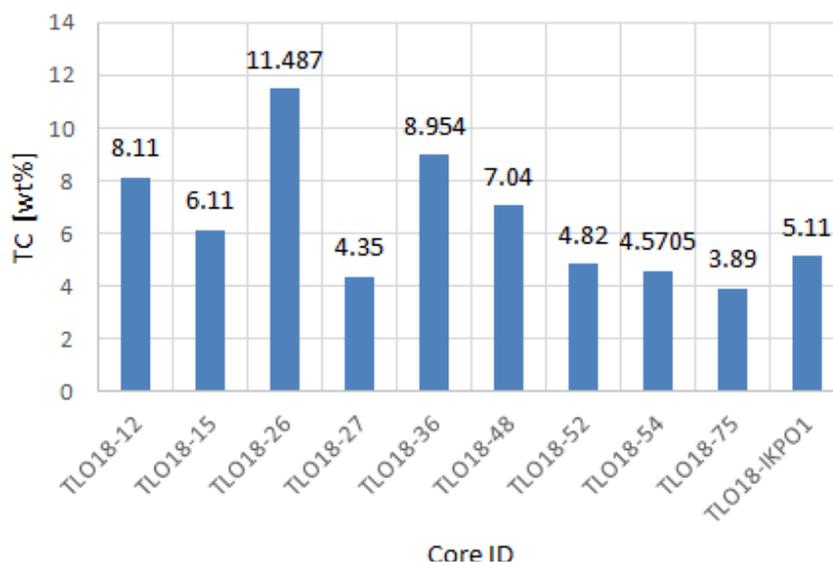


Figure 10: Total carbon of all cores

4.2.3 Total organic carbon

The TOC values are in a very similar range (from 3.67 wt% to 11.337 wt %) as the TC values and correspond very well. The highest values were also measured in TLO18-12, -26 and -36. Otherwise the lowest were measured in TLO18-27, -54 and -75 with 4.35 wt%, 4.5705 wt% and 3.89 wt% again. TLO18-IKPO1 contains more TOC (5.23 wt%) than TC (5.11 wt%) (**Figure 11**).

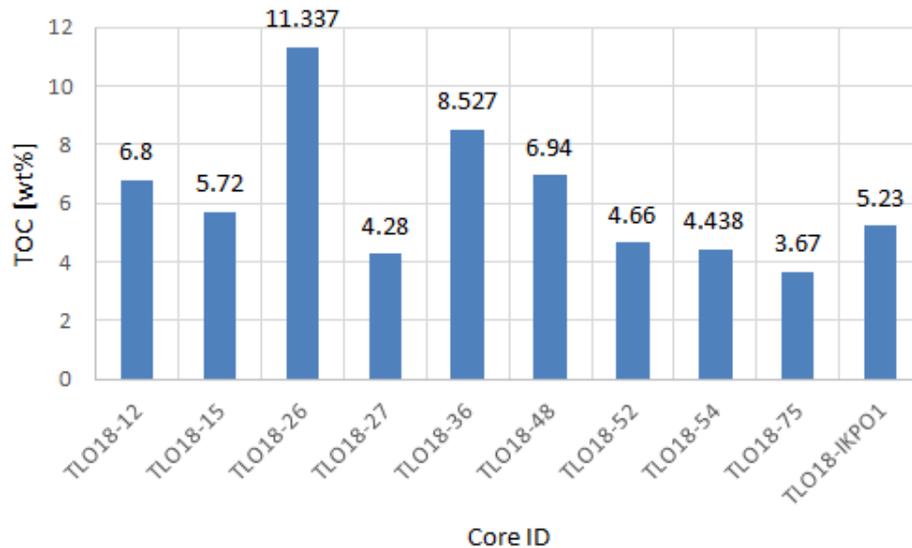


Figure 11: Total organic carbon in all cores

4.2.4 Stable carbon isotopes

The $\delta^{13}\text{C}$ values reach from -27.16 ‰VPDB to -29.52 ‰VPDB. The lowest values were measured in TLO18-IKPO1, -26, -36 and -12 which also describe the highest ratios. They spread between -28.68 ‰VPDB to -29.52 ‰VPDB. The highest values were measured in TLO18-27 and -75 with -28.3 ‰VPDB and -27.16 ‰VPDB which is a lower ratio of stable carbon isotopes (**Figure 12**).

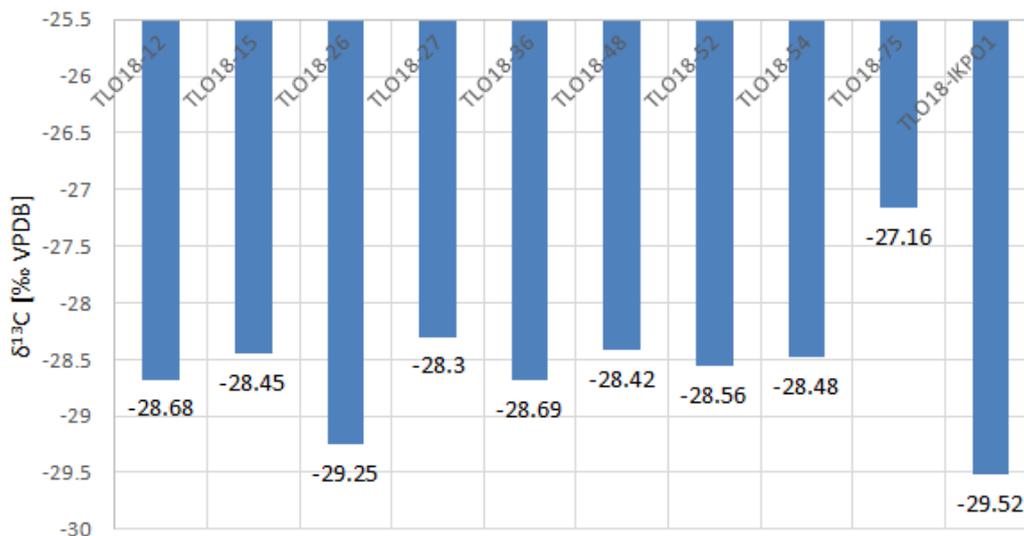


Figure 12: Stable carbon isotopes ratio of all cores

4.3 GIS analytics

4.3.1 Lake size

The digitalisations of lake sizes show spatial variabilities over the years 1986 to 2018. (**Table 3, Figure 13, 14**) The most dynamics were visible on the lakes 27, 12 and 36. They register a size change of 7.18% to 5.82 %. The lowest changes were calculated on lake 75 with 0.25 % and lake 54 with 3.89 %. Every lake, except lake 27, register a positive lake size growth. The lake 27 shows a negative lake growth and implies a drainage.

Table 3: Lake dynamics (size changes) from 1986 to 2018

Lake ID	Lake size 1986 [km ²]	Lake size 2018 [km ²]	Dynamic [km ²]	Dynamic [%]
12	9,844	10,485	0,641	6,11
15	9,572	10,018	0,446	4,45
26	6,449	6,723	0,274	4,08
27	4,028	3,758	-0,270	-7,18
36	5,535	5,877	0,342	5,82
48	3,809	3,987	0,178	4,46
52	18,524	19,505	0,981	5,03
54	16,727	17,404	0,677	3,89
75	13,035	13,068	0,033	0,25
IKPO1	0,652*	0,689	0,037	5,37

*Lake size was digitalised with Landsat 5 image from 25th July 2006

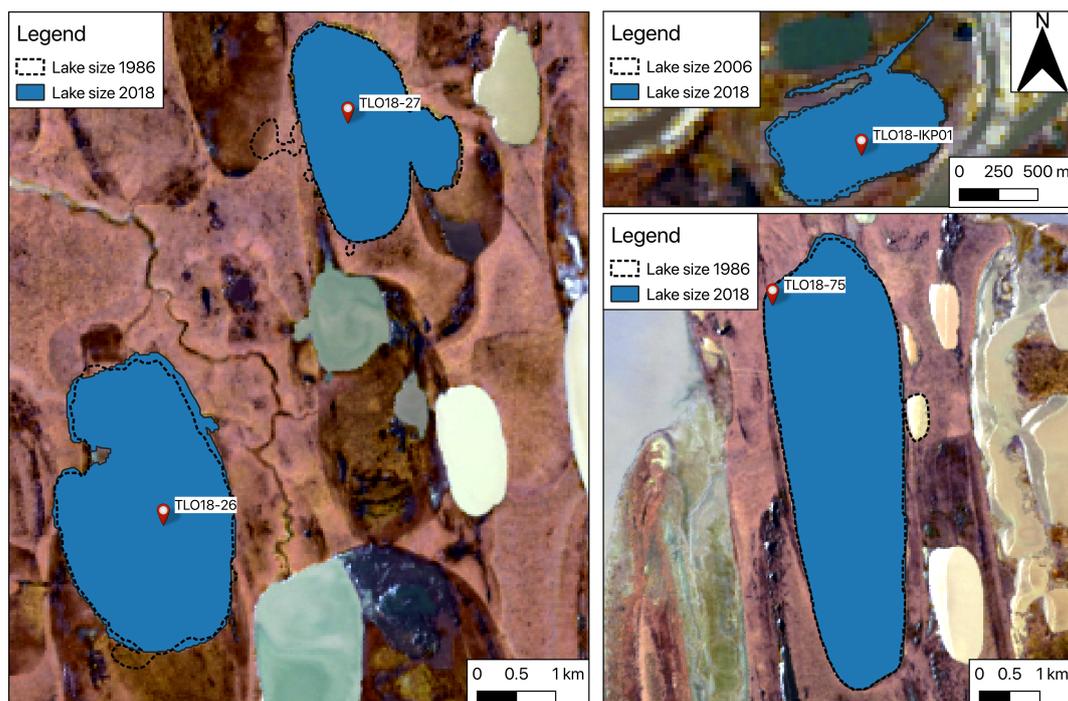


Figure 13: Lake size dynamics of Lakes 26, 27, IKPO1 and 75 (Backmap from Landsat 8 satellite imagery provided by USGS Data Center)

5. Discussion

5.1 Interpretation of organic matter characteristics

Most of the made visual characterisation corresponds with the measured TOC content as well. As TLO18-26, -36, -48 and -IKPO1 were classified as organic dominated, they also show the highest TOC contents. On the other hand, it is shown that TLO18-12 contains even more total organic carbon than TLO18-IKPO1 which was not suspected while describing the core because it was characterised as minerogenic dominated (**Table 2**).

As the results show, the carbon content of TC and TOC measurements are corresponding. The TOC in TLO18-IKPO1 is higher than the measured TC values. This represents that there is barely any inorganic carbon in this core. Higher inorganic carbon content would be shown as remains of shells or crustacean which is not a significant result of this study.

As TN values provide the bio productivity of each core. The cores TLO18-12, -26, -36 and -48 show values >0.4 wt% and >8.2 wt% TOC which indicates a higher bioproductivity (Lenz et al., 2013).

The C/N ratio shows shows the degree of decomposition (**Figure 16**), while the ratio of C/N and $\delta^{13}\text{C}$ provide the discrimination of organic sources or vegetation origin (**Figure 17**, Lenz et al., 2016).

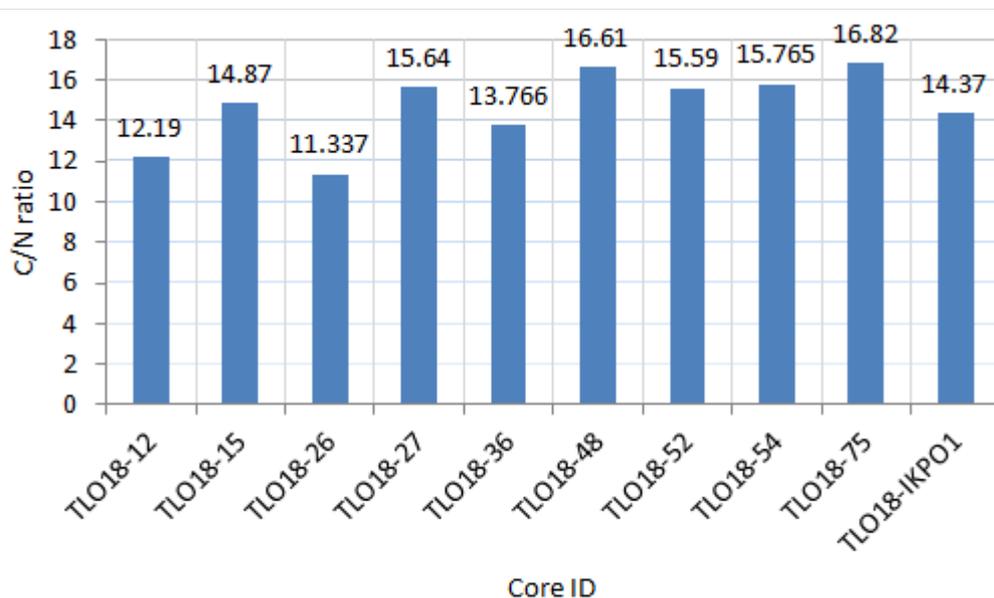


Figure 16: C/N ratio of all cores

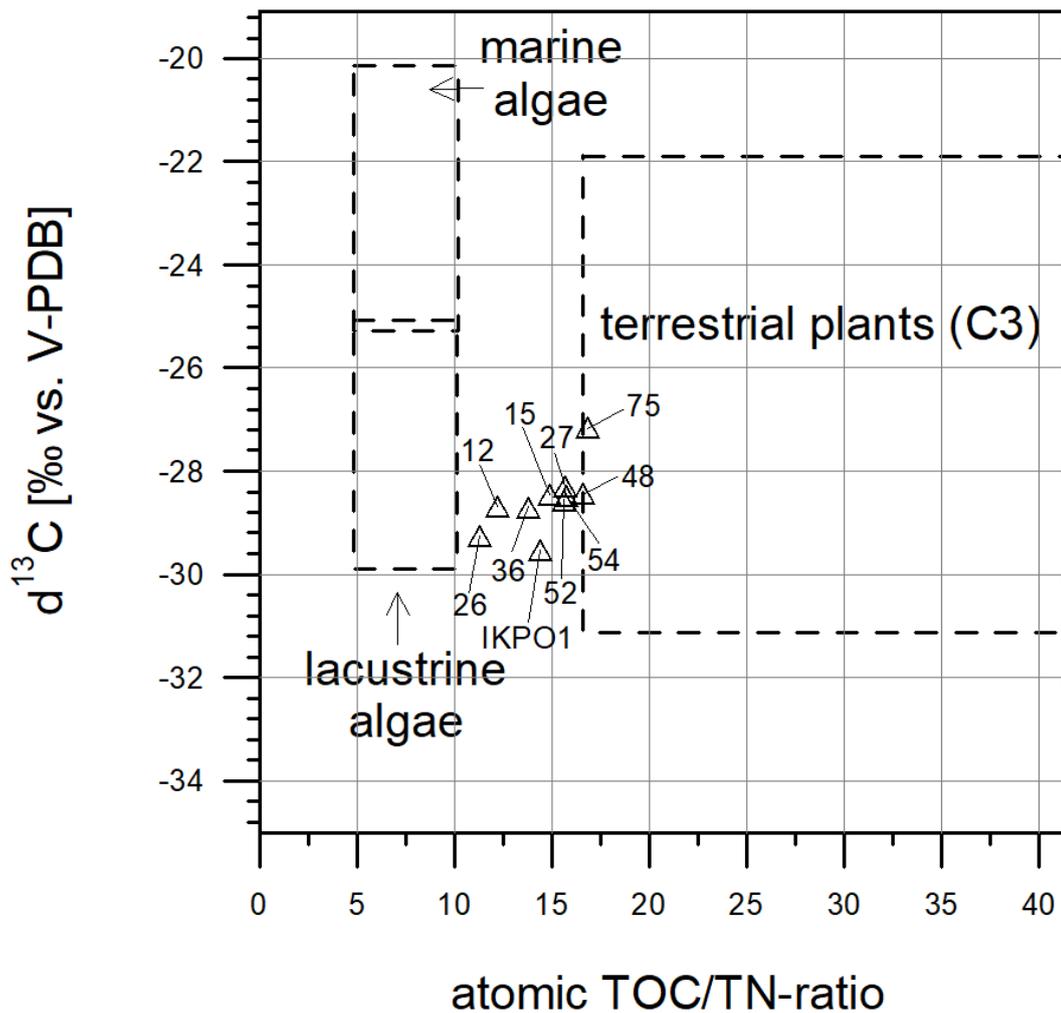


Figure 17: C/N ratio and $\delta^{13}\text{C}$ values of all cores

5.2 GIS analytics

The measured dynamic of the lakes was expected to occur mostly as a positive growth, which also happened as expected, because of the increasing ground temperatures. (Romanovski et al., 2010) On the other side, a lot of thermokarst lakes in this area lean to drain, thus there is resulting a high thermo-erosion in the ground (Jones et al., 2015).

It is also interesting to have a look on the growth direction because there is no growth in all directions. Lakes 12, 15, 26, 36, 48, 52, 54 and 75 all tend to expand in northern directions, thus they stretch to the coastal area. (Figure 13 and 14, Arp et al., 2011)

While studying the grounded and floating ice lakes, there showed a spatial distribution of every lake type. Grounded ice lakes seems rather to appear in the east part of the study area, while floating ice lakes seems to be rather in the centre and west of the study area.

5.3 Work Hypotheses

5.3.1 First work hypothesis

The analytical results were used to show how different lake parameters relate to each other. First work hypotheses express an assumed relation between stable lakes, more decomposed organic and rather aquatic plant remains in comparison to dynamic lakes which contain less decomposed organics and rather terrestrial plants. To prove this hypothesis, it is necessary to prove if there are lakes which show these assumed parameters at the same time. While Lakes 26, 48, 54 and 75 are classified as stable lakes. The lakes 12, 15, 27, 36, 52 and IKPO1 are categorised as dynamic lakes. To examine the rate of decomposed organics it is necessary to use the C/N ratio. It classifies the stable lakes 48, 54 and 75 as the most containing decomposed organic. However, the fourth stable lake 26 is classified containing the lowest decomposed organic material which does not fit to the assumed hypothesis and which is not suitable for further proving. To verify the vegetation origin, it is necessary to show a relation between $\delta^{13}C$ values and C/N ratio. It shows that the stable lakes 48, 54 and 75 contain more decomposed organic and do not contain aquatic plants. The cores TLO18-48 and -75 contain terrestrial plants which shows clearly the opposite of the assumed hypothesis. Even core TLO18-54 contains rather terrestrial plants. Otherwise, the other stable lake with less decomposed organic contains lacustrine algae. As a result of this, it could be said that stable lakes contain rather decomposed organic but it can not be proved that they also contain rather aquatic material which seems to depend on its decomposing level as well.

The dynamic lakes 12, 15, 27, 36, 52 and IKPO1 are assumed to contain less decomposed organic. This is accurate for lakes 12, 36 and IKPO1. Otherwise, TLO18-15, -27 and -52 show higher decomposition values which is then not suitable for further analytics in relation to first hypothesis. TLO18-12 contains rather lacustrine algae which is the opposite of the supposed hypothesis. TLO18-36 and IKPO1 contain rather terrestrial plants which is accurate in relation to the first hypothesis. The other dynamic lakes, which contain rather decomposed organics, contain rather terrestrial plants as well. All in all, the first hypothesis can not be confirmed because there are some contradictions and only two lakes are accurate.

Table 4: Summarised results for the first work hypothesis, Dynamic, C/N ratio and Vegetation origin

Core ID	Dynamic – Lake size change [%]	C/N ratio (TOC/TN)	Vegetation origin
TLO18-12	6,11349547	12,19	rather lacustrine algae
TLO18-15	4,451986424	14,87	rather terrestrial
TLO18-26	4,075561505	11,337	rather lacustrine algae
TLO18-27	-7,184672698	15,64	rather terrestrial
TLO18-36	5,819295559	13,766	rather terrestrial
TLO18-48	4,464509656	16,61	terrestrial
TLO18-52	5,029479621	15,59	rather terrestrial
TLO18-54	3,889910365	15,765	rather terrestrial
TLO18-75	0,252525253	16,82	terrestrial
TLO18-IKPO1	5,370101597	14,37	rather terrestrial
Classification			
	Stable (± 0.1 - ± 2.9)	Low (11,3-12,5)	
	Rather stable (± 3 - ± 4.4)	Rather low (12,51-14,5)	
	Rather dynamic (± 4.5 - $\pm 5,9$)	Rather high (14,51-15,5)	
	Dynamic (± 6 - ± 7.2)	High (15,51-17)	

5.3.2 Second work hypothesis

The second work hypotheses stated a relation between smaller, shallower and grounded ice lakes and their lower organic matter input in comparison to larger, deeper and floating ice lakes with higher organic matter input. Once again it was necessary to prove, if the parameters would correspond with each other. As small lakes were classified lake 26, 27, 36, 48 and IKPO1. Lake 12 15 52, 54 and 75 were categorised as larger lakes. The smaller lakes are assumed to be shallower which is accurate for lake 27, 36 and 48. Nevertheless, lake 26 and IKPO1 are categorised as deeper lakes which causes a discrimination of these lakes for further analytics in relation to the ice lake type and their organic matter input. The smaller lakes with shallower water depths that also are classified as grounded ice lakes are lake 27 and lake 48, hence lake 36 is a floating ice lake which makes it not accurate in relation to the organic matter input rate. Lake 27 exhibits a lower organic matter input. However, lake 48 shows a higher organic matter input which contradicts the second work hypothesis. So one of five possible lakes is accurate in relation to the second work hypothesis so far.

The second work hypothesis needs to be proved by the opposite parameters as well.

To the larger lakes with deeper water depths belong lake 12 and 15. Lake 52, 54 and 75 are larger lakes with shallower water depths which makes them not suitable for the second work hypothesis either. If lake 12 and 15 are proved in relation to their ice lake type, they show assumed results because they are classified as floating ice lakes. All in all, the second hypothesis is accurate for three out of ten expected lakes which shows that the opposite of the hypothesis is more accurate. So according to the results, it could be expected that grounded ice lakes still exhibit less organic matter input and floating ice lakes rather more organic matter input. The issue with this hypothesis is the missing relation between lakes sizes and water depths because there is no clear declaration as there are two large and deep lakes, three small and shallow lakes, two small and deep lakes and three large and shallow lakes.

Table 5: Summarised results for the second work hypothesis, lake size, lake, depth, ice lake type, TOC

core ID	Lake size 2018 [km ²]	Lake Depth [cm]	Ice lake type [Arp et al., 2012]	TOC [wt %]
TLO18-12	10.485	250	floating	6.8
TLO18-15	10.018	220	floating	5.72
TLO18-26	6.723	230	floating	11.337
TLO18-27	3.758	64	grounded	4.28
TLO18-36	5.877	200	floating*	8.527
TLO18-48	3.987	160	grounded	6.94
TLO18-52	19.505	200	floating	4.66
TLO18-54	17.404	100	grounded	4.438
TLO18-75	13.068	100	grounded	3.67
TLO18-IKPO1	0.689	350	-	5.23
Classification				
Low	0-4	0-100		0-4
Rather low	4.1-8	101-200		4.1-5.0
Rather high	8.1-12	201-250		5.1-8.0
High	12.1-18	251-400		8.1-12

*obtained from Arp et al., 2011

5.3.3 Third work hypothesis

The third work hypothesis is assuming that lakes with higher salinity contain less organic matter than lakes with lower salinity. Lakes with a higher salinity are lake 48, 75 and IKPO1. They are influenced by marine waters a lot regarding to their degree of salinity because marine waters indicate an electrical conductivity of >2000 $\mu\text{S}/\text{cm}$. (Clean Water Team, 2004) Regarding to their salinity and the less organic matter input is only lake 75 consisting, hence lake 48 and IKPO1 contain high amounts of TOC which shows that only one of three lakes verify the one side of the third hypothesis. Lakes with low salinity and high organic matter input are lake 12, 15, 26 and 36. This shows that the third work hypothesis is suitable for five out of ten lakes.

According to this hypothesis, it could be also expected that higher salinity means also a higher organic input and a lower salinity means also a lower organic matter input. Because of the 50 % of accuracy this thesis can not be proved correctly.

Table 6: Summarised results for the third work hypothesis, electric conductivity, TOC

Core ID	Electrical conductivity [$\mu\text{S}/\text{cm}$]	TOC [wt%]
TLO18-12	400	6.8
TLO18-15	485	5.72
TLO18-26	570	11.337
TLO18-27	422	4.28
TLO18-36	507	8.527
TLO18-48	10380	6.94
TLO18-52	168	4.66
TLO18-54	143	4.438
TLO18-75	7000	3.67
TLO18- IKPO1	8020	5.23
Classification		
Low	0-300	0-4
Rather low	301-1000	4.1-5.0
Rather high	1001-4500	5.1-8.0
High	4501-11000	8.1-12

*thick black lines mark consisting lakes

6. Outlook

As the classification and relation between each parameter was made qualitatively, there is also the possibility of validating the work hypotheses with a Principal Component Analysis which can statistically show the relation of every parameter with each other.

This study is not based on specific radiocarbon dating which makes it difficult to show when the dynamics specifically happened. On the other side, the time parameter was not necessary for this study because of using the average of the first 10 cm of the top of each core which caused a discrimination of the time parameter.

The bio productivity could have been compared with pH values to see if these parameters relate to each other to create a more biological focus.

7. Conclusion

Within this study the following results could have been conducted:

- Thermokarst lakes are very sensitive to climate change which shows in positive growth rates because of abrupt thawing. (**Table 3**)
- According to stable and dynamic lakes, degree of decomposition and aquatic and terrestrial there could no relation be proved.
- Grounded and floating ice lakes seems to have a relation with the organic input matter. Otherwise, there is no relation with lake depths and lake sizes.
- Five of ten lakes seem to have a relation of organic matter input and salinity.

8. Attachment

Arp, C. D., B. M. Jones, F. E. Urban, G. Grosse (2011): Hydrogeomorphic processes of thermokarst lakes with grounded-ice and floating-ice regimes on the Arctic coastal plain, Alaska, HYDROLOGICAL PROCESSES, DOI: 10.1002/hyp.8019

Arp, Ch. D., Jones, B. M., Lu, Z., Whitman, M. S. (2012): Shifting balance of thermokarst lake ice regimes across the Arctic Coastal Plain of northern Alaska, Geophysical research letters, DOI:10.1029/2012GL052518

Clean Water Team (CWT) (2004): Electrical conductivity/salinity Fact Sheet, FS- 3.1.3.0(EC). in: The Clean Water Team Guidance Compendium for Watershed Monitoring and Assessment, Version 2.0. Division of Water Quality, California State Water Resources Control Board (SWRCB), Sacramento, CA

Grosse, G., Jones B., and Arp C. (2013): Thermokarst Lakes, Drainage, and Drained Basins. In: John F. Shroder (Editor-in-chief), Giardino, R., and Harbor, J. (Volume Editors), Treatise on Geomorphology, Vol 8, Glacial and Periglacial Geomorphology, San Diego: Academic Press; 2013. p. 325-353.

Institut Dr. Fled (): Die Bestimmung eines Chemischen Index zur Ermittlung der Gewässergüteklasse von Fließgewässern, available at:
https://www.chf.de/eduthek/chemischer-index/Chemischer_Index.pdf

Jones, B. M., Arp, C. D. (2015): Observing a catastrophic thermokarst lake drainage in Northern Alaska, Permafrost and Periglacial Processes, DOI: 10.1002/ppp.1842

Jorgenson, J., Yoshikaw, T., Kanevskiy, K., Shur, M., Romanovsky, V. E., Marchenko, V., Grosse, G. Brown, B. M. Jones (2008): Permafrost characteristics of Alaska, Institute of Northern Engineering, University of Alaska, Fairbanks, AK, Map and text

Lenz, J., Fritz, M., Schirrmeister, L., Lantuit, H., Wooller, M. J., Pollard, W. H., Wetterich, S. (2013): Periglacial landscape dynamics in the western Canadian Arctic : Results from a thermokarst lake record on a push moraine (Herschel Island, Yukon Territory), ELSEVIER

Lenz, J., Jones, B. M., Wetterich, S., Tjallingii, R., Fritz, M., Arp, C. D., Grosse, G. (2016): Impacts of shore expansion and catchment characteristics on lacustrine thermokarst records in permafrost lowlands, Alaska Arctic Coastal Plain, *Arktos* 2:25.

Mars, J. C., Houseknecht, D. W., (2007): Quantitative remote sensing study indicates doubling of coastal erosion rate in past 50 yr along a segment of the Arctic coast of Alaska, *GeoScienceWorld*, <https://doi.org/10.1130/G23672A.1>

Munsell, (1994): Soil, color charts, revised edition, Nova York, MacBeth Division of Kollmorgan Instruments Corporation

NOAA: Weather data from Barrow, Alaska (w.Pos. United States of America) 1987-2016 via <http://climatecharts.net>

Nowacki, G., Spencer, P., Fleming, M., Brock, T., Jorgenson, J. (2002): Unified ecoregions of Alaska, U.S. Geological Survey Open File Report

Overland, J., Hanna, E., Hanssen-Bauer, I., Kim, S.-J., Walsh, J., Wang, M., Bhatt, U. S., Thomann, R. L. (2017): Surface air temperature, available on: <http://www.arctic.noaa.gov/Report-Card/Report-Card-2017>

Pithan, F., Mauritsen, T. (2014): Arctic amplification dominated by temperature feedback in contemporary climate models, *nature geoscience*, <https://doi.org/10.1038/ngeo2071>

Romanovsky, V. E., Smith, S. L., Christiansen H. H. (2010): Permafrost Thermal State in the Polar Northern Hemisphere during the International Polar Year 2007 2009: a Synthesis, Permafrost and Periglacial Processes, DOI: 10.1002/ppp.689

Shulski, M, Wendler, G. (2007): The climate of Alaska, University of Alaska Press

Walter Anthony, K. M. (2014): A shift of thermokarst lakes from carbon sources to sinks during the Holocene epoch, Nature 511.7510 (2014): 452.

Walter Anthony, K. M., Schneider von Deimling, T., Nitze, I., Froking, S., Emond, A., Daanen, R., Anthony, P., Lindgren, P., Jones, B., Grosse, G. (2018): 21st-century modeled permafrost carbon emissions accelerated by abrupt thaw beneath lakes, nature communications, DOI: 10.1038/s41467-018-05738-9

Walter, K. M., E. Edwards. G. Grosse, S. A. Zimov, F. S. Chapin III (2007): Thermokarst Lakes as a Source of Atmospheric CH₄ During the Last Deglaciation, Science

Woodward, A., Beever, E. A. (2011): Conceptual ecological to support detection of ecological change on Alaska National Wildlife Refuges, U.S. Geological Survey Open-File Report 2011

Satellite imageries used in this study:

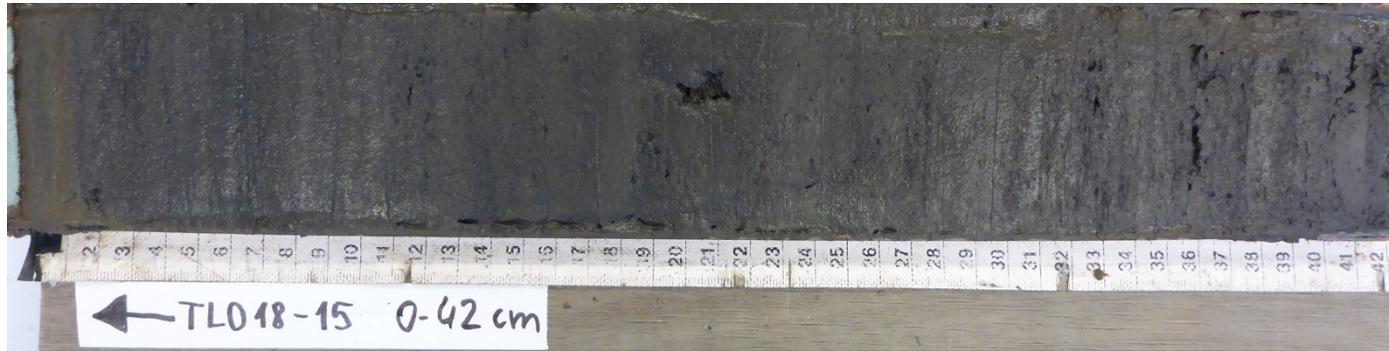
Satellite	Imagery ID	Date (DD.MM.YYYY)	Cloud cover [%]	Source
Landsat 5	LANDSAT/LT05/C01/T1_SR/LT05_076010_19860812 (11 bands)	12.08.1986	31	U.S. Geological Survey Data Center
Landsat 5	LANDSAT/LT05/C01/T1_SR/LT05_077010_20060725 (11 bands)	25.07.2006	19	U.S. Geological Survey Data Center
Landsat 8	LANDSAT/LC08/C01/T1_SR/LC08_078010_20180919 (12 bands)	19.09.2018	10.44	U.S. Geological Survey Data Center

9.Appendix
Tables of data
TLO18-12



Sample Depth [cm]	Short description	Water content [%]	TN [wt%]	TC [wt%]	TOC [wt%]	TOC/N	δ 13C
2.0	fU to T, minberogenic	62.36	0.549	7.734	6.730	12.2533634	-28.78
5.0	fU to T, minberogenic	66.38	0.619	9.232	7.432	12.00320911	-28.79
9.0	fU to T, minberogenic	49.08	0.508	7.364	6.250	12.3059831	-28.47
12.0	fU to T, minberogenic	52.91	0.521	7.500	5.547	10.65516955	-28.52
15.0	fU to T, organic rich	50.15	0.526	8.191	7.551	14.36941754	-28.52
19.0	fU to T, organic rich	50.61	0.501	8.059	7.406	14.7920533	-28.59
24.0	fU to T, organic rich	56.85	0.604	10.317	8.574	14.18707863	-28.49
28.0	fU to T, organic rich	49.39	0.429	7.351	6.234	14.5242568	-28.43
32.0	fU to T, organic rich	50.33	0.491	8.394	6.776	13.79993878	-28.30
36.0	fU to T, organic rich	38.58	0.240	4.245	4.125	17.19313274	-28.12
39.0	fU to T, organic rich	43.88	0.328	5.828	5.524	16.83924283	-28.08
Used for study							
Sample Depth [cm]	Short description	Water content [%]	TN [wt%]	TC [wt%]	TOC [wt%]	TOC/N	δ 13C
2.0	fU to T, minberogenic	62.36	0.549	7.734	6.730	12.2533634	-28.78
5.0	fU to T, minberogenic	66.38	0.619	9.232	7.432	12.00320911	-28.79
9.0	fU to T, minberogenic	49.08	0.508	7.364	6.250	12.3059831	-28.47
AVERAGE		59.27	0.56	8.11	6.80	12.19	-28.68

TLO18-15



Sample Depth [cm]	Short description	Water content [%]	TN [wt%]	TC [wt%]	TOC [wt%]	TOC/N	$\delta^{13}C$
1.0	fS, minerogenic	43.30	0.215	3.405	2.955	13.74323534	-28.15
4.0	gU to fS, minerogenic	61.25	0.458	7.333	6.954	15.1842489	-28.62
8.0	gU to fS, minerogenic	59.34	0.462	7.589	7.247	15.6868273	-28.57
13.0	gU to fS, minerogenic	57.06	0.466	7.369	6.966	14.94858807	-28.56
18.0	gU to fS, minerogenic	55.16	0.425	7.055	6.615	15.56543631	-28.81
23.0	gU to fS, minerogenic	59.51	0.506	8.033	7.761	15.33739934	-28.89
28.0	gU to fS, minerogenic	64.04	0.593	9.724	9.219	15.54604284	-28.81
33.0	gU to fS, minerogenic but organic remains visible	50.21	0.304	5.59	5.307	17.45833378	-28.22
36.0	gU to fS, minerogenic but organic remains visible	58.77	0.427	7.717	7.332	17.17006127	-28.38
38.5	gU to fS, minerogenic but organic remains visible	30.60	0.2	3.268	2.936	14.6798259	-28.19
41.0	T to fU	28.71	0.276	4.538	4.289	15.54085293	-28.16
Used for study							
Sample Depth [cm]	Short description	Water content [%]	TN [wt%]	TC [wt%]	TOC [wt%]	TOC/N	$\delta^{13}C$
1.0	fS, minerogenic	43.30	0.215	3.405	2.955	13.74323534	-28.15
4.0	gU to fS, minerogenic	61.25	0.458	7.333	6.954	15.1842489	-28.62
8.0	gU to fS, minerogenic	59.34	0.462	7.589	7.247	15.6868273	-28.57
AVERAGE		54.63	0.38	6.11	5.72	14.87	-28.45

TLO18-26



Sample Depth [cm]	Short description	Water content [%]	TN [wt%]	TC [wt%]	TOC [wt%]	TOC/N	δ 13C
1.0	fU, well decomposed organic mud with poorly decomposed remains	69.77	0.731	12.111	12.077	16.52097832	-29.32
5.0	"	63.54	0.647	11.212	10.819	16.72174094	-29.15
10.0	"	67.53	0.644	11.138	11.114	17.2576349	-29.27
13.0	"	65.57	0.734	13.168	13.435	18.30418233	-28.91
17.0	"	64.90	0.618	10.546	10.263	16.60610866	-29.21
22.0	"	62.92	0.674	11.842	11.736	17.41189348	-29.05
27.0	"	61.68	0.665	11.399	11.250	16.9165726	-29.04
32.0	"	65.43	0.671	11.327	11.052	16.47140301	-29.40
37.0	"	52.13	0.638	10.38	9.968	15.6231928	-28.87
42.0	"	57.84	0.53	9.166	8.840	16.67881012	-28.88
47.0	"	64.00	0.676	11.686	11.076	16.38494723	-29.03
52.0	"	62.47	0.661	11.604	11.153	16.87258852	-29.12
56.0	"	63.86	0.771	13.645	13.921	18.05618363	-29.20
61.0	"	49.54	0.581	10.191	9.904	17.04683895	-29.20
Used for study							
Sample Depth [cm]	Short description	Water content [%]	TN [wt%]	TC [wt%]	TOC [wt%]	TOC/N	δ 13C
1.0	"	69.77	0.731	12.111	12.077	16.52097832	-29.32
5.0	"	63.54	0.647	11.212	10.819	16.72174094	-29.15
10.0	"	67.53	0.644	11.138	11.114	17.2576349	-29.27
AVERAGE		66.95	0.674	11.487	11.337	16.83345139	-29,25

TLO18-27



Sample Depth [cm]	Short description	Water content [%]	TN [wt%]	TC [wt%]	TOC [wt%]	TOC/N	$\delta^{13}\text{C}$
1.0	mU, minerogenic	35.74	0.272	4.282	4.225	15.53403104	-28.31
5.0	mU, minerogenic	31.85	0.269	4.325	4.249	15.79426567	-28.29
9.0	mU, minerogenic	30.16	0.281	4.455	4.380	15.58606905	-28.29
12.5	T to fU	29.59	0.312	5.203	4.997	16.01502758	-28.43
Used for study							
Sample Depth [cm]	Short description	Water content [%]	TN [wt%]	TC [wt%]	TOC [wt%]	TOC/N	$\delta^{13}\text{C}$
1.0	mU, minerogenic	35.74	0.272	4.282	4.225	15.53403104	-28.31
5.0	mU, minerogenic	31.85	0.269	4.325	4.249	15.79426567	-28.29
9.0	mU, minerogenic	30.16	0.281	4.455	4.380	15.58606905	-28.29
AVERAGE		32.58	0.27	4.35	4.28	15.64	-28.30

TLO18-36



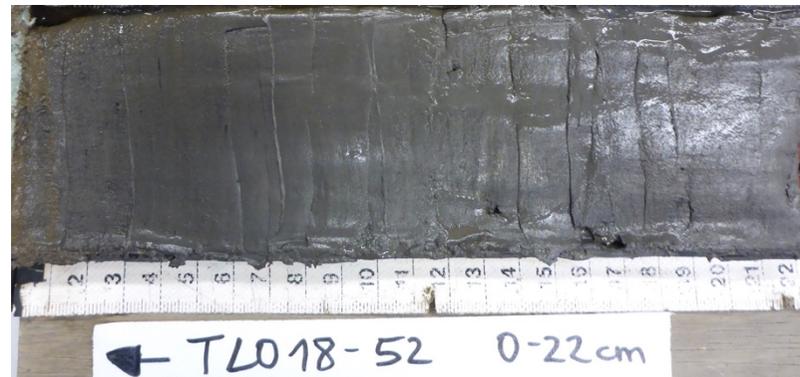
Sample Depth [cm]	Short description	Water content [%]	TN [wt%]	TC [wt%]	TOC [wt%]	TOC/N	$\delta^{13}C$
1.0	fU, organic mud, well decomposed	67.86	0.578	8.576	8.240	14.26830551	-28.72
5.0	fU, organic mud, well decomposed	68.16	0.688	9.392	8.839	12.84150942	-28.63
9.0	fU, organic mud, well decomposed	61.42	0.599	8.894	8.503	14.18910281	-28.72
12.0	fU, coarser organic remains	70.10	0.620	9.289	8.817	14.21439739	-28.73
16.0	fU, coarser organic remains	64.91	0.697	11.101	10.997	15.76898626	-28.57
20.0	gU, coarser minerogenic, organic remains	57.21	0.594	9.149	8.898	14.97059075	-28.79
25.0	gU, coarser minerogenic, organic remains	52.15	0.526	8.669	8.370	15.92654625	-28.66
29.0	gU, coarser minerogenic, organic remains	50.34	0.437	7.605	7.423	16.97845531	-28.23
33.0	gU, coarser minerogenic, organic remains	38.69	0.267	4.608	4.526	16.97900538	-28.28
Used for study							
Sample Depth [cm]	Short description	Water content [%]	TN [wt%]	TC [wt%]	TOC [wt%]	TOC/N	$\delta^{13}C$
1.0	fU, organic mud, well decomposed	67.86	0.578	8.576	8.240	14.26830551	-28.72
5.0	fU, organic mud, well decomposed	68.16	0.688	9.392	8.839	12.84150942	-28.63
9.0	fU, organic mud, well decomposed	61.42	0.599	8.894	8.503	14.18910281	-28.72
AVERAGE		65.82	0.622	8.954	8.527	13.76630591	-28.69

TLO18-48



Sample Depth [cm]	Short description	Water content [%]	TN [wt%]	TC [wt%]	TOC [wt%]	TOC/N	δ 13C
1.0	T to fU	62.77	0.387	6.387	6.204	16.03036826	-28.44
4.0	T to fU	59.54	0.442	7.669	7.744	17.52004472	-28.36
7.0	fU, no organic remains, well decomposed?	57.51	0.422	7.056	6.866	16.27021152	-28.45
12.0	fU, no organic remains, well decomposed?	60.59	0.54	9.811	9.437	17.47571274	-28.44
15.0	fU, no organic remains, well decomposed?	60.02	0.471	8.799	8.690	18.44952415	-28.32
17.0	organic rich, macro remains	77.34	0.731	14.813	14.741	20.1649144	-28.75
20.0	gU, minerogenic, moderately decomposed organics	50.43	0.385	7.111	6.750	17.53224893	
24.0	gU, minerogenic, moderately decomposed organics	50.94	0.5	8.863	8.462	16.92426586	-28.50
28.0	gU, minerogenic, moderately decomposed organics	57.30	0.631	11.299	11.027	17.4762277	-28.56
31.5	T	31.47	0.182	3.612	3.511	19.28996778	-27.86
Used for study							
Sample Depth [cm]	Short description	Water content [%]	TN [wt%]	TC [wt%]	TOC [wt%]	TOC/N	δ 13C
1.0	T to fU	62.77	0.387	6.387	6.204	16.03036826	-28.44
4.0	T to fU	59.54	0.442	7.669	7.744	17.52004472	-28.36
7.0	fU, no organic remains, well decomposed?	57.51	0.422	7.056	6.866	16.27021152	-28.45
AVERAGE		59.94	0.42	7.04	6.94	16.61	-28.42

TLO18-52



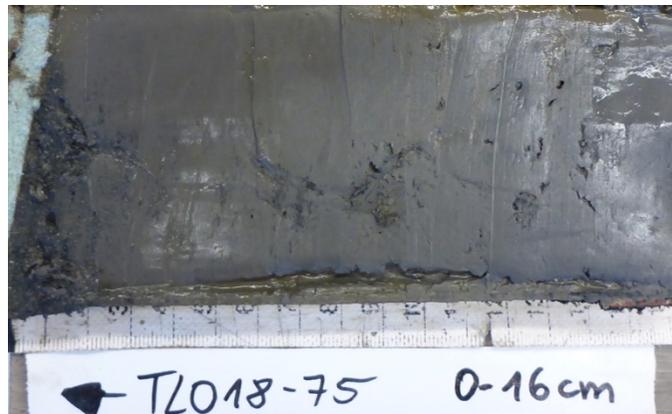
Sample Depth [cm]	Short description	Water content [%]	TN [wt%]	TC [wt%]	TOC [wt%]	TOC/N	$\delta^{13}C$
1.0	mU, minerogenic	41.59	0.283	4.537	4.314	15.24319666	-28.52
4.0	mU, minerogenic	32.27	0.284	4.62	4.441	15.63881485	-28.39
9.0	mU, minerogenic	32.93	0.328	5.306	5.213	15.893886	-28.77
13.0	mU, minerogenic	34.29	0.303	5.238	5.262	17.36779732	-28.73
17.0	mU, minerogenic	33.38	0.319	5.502	5.498	17.23552647	-28.72
21.0	mU, minerogenic	35.27	0.39	6.814	6.553	16.80354644	-28.79
Used for study							
Sample Depth [cm]	Short description	Water content [%]	TN [wt%]	TC [wt%]	TOC [wt%]	TOC/N	$\delta^{13}C$
1.0	mU, minerogenic	41.59	0.283	4.537	4.314	15.24319666	-28.52
4.0	mU, minerogenic	32.27	0.284	4.62	4.441	15.63881485	-28.39
9.0	mU, minerogenic	32.93	0.328	5.306	5.213	15.893886	-28.77
AVERAGE		35.60	0.30	4.82	4.66	15.59	-28.56

TLO18-54



Sample Depth [cm]	Short description	Water content [%]	TN [wt%]	TC [wt%]	TOC [wt%]	TOC/N	δ 13C
1.0	fU, no visible organic remains	39.57	0.3	4.964	4,829	16.09667699	-28.64
3.0	T, no visible organic remains	32.68	0.285	4.835	4,663	16.36065349	-28.52
6.0	T, no visible organic remains	30.36	0.276	4.344	4,158	15.06482432	-28.40
9.0	T, no visible organic remains	28.96	0.264	4.139	4,102	15.53956126	-28.37
11.5	T, no visible organic remains	28.57	0.278	4.264	4,125	14.83924183	-28.41
Used for study							
Sample Depth [cm]	Short description	Water content [%]	TN [wt%]	TC [wt%]	TOC [wt%]	TOC/N	δ 13C
1.0	fU, no visible organic remains	39.57	0.3	4.964	4,829	16.09667699	-28.64
3.0	T, no visible organic remains	32.68	0.285	4.835	4,663	16.36065349	-28.52
6.0	T, no visible organic remains	30.36	0.276	4.344	4,158	15.06482432	-28.40
9.0	T, no visible organic remains	28.96	0.264	4.139	4,102	15.53956126	-28.37
AVERAGE		32.89	0.28125	4.5705	4,438031256	15.76542901	-28.48295253

TLO18-75



Sample Depth [cm]	Short description	Water content [%]	TN [wt%]	TC [wt%]	TOC [wt%]	TOC/N	δ 13C
1.5	T to fU, living vegetation on top	47.32	0.311	6.213	6.233	20.04109318	-28.05
3.5	T to fU	26.53	0.162	2.729	2.401	14.82385103	-26.74
7.0	T to fU	24.79	0.153	2.732	2.387	15.60116905	-26.70
11.0	T to fU	23.97	0.172	2.686	2.359	13.71277973	-26.72
15.0	T to fU	26.37	0.175	3.045	2.648	15.13289315	-26.68
Used for study							
Sample Depth [cm]	Short description	Water content [%]	TN [wt%]	TC [wt%]	TOC [wt%]	TOC/N	δ 13C
1.5	T to fU, living vegetation on top	47.32	0.311	6.213	6.233	20.04109318	-28.05
3.5	T to fU	26.53	0.162	2.729	2.401	14.82385103	-26.74
7.0	T to fU	24.79	0.153	2.732	2.387	15.60116905	-26.70
AVERAGE		32.88	0.21	3.89	3.67	16.82	-27.16

TLO18-IKPO1



Sample Depth [cm]	Short description	Water content [%]	TN [wt%]	TC [wt%]	TOC [wt%]	TOC/N	$\delta^{13}C$
1.0	mU, minerogenic	49.70	0.255	3.48	3.301	12.94521023	-28.38
3.0	mU to gU, finely decomposed organic	69.81	0.5	???	5.812	11.62400055	-29.26
6.0	mU to gU, finely decomposed organic	72.22	0.379	5.721	5.760	15.19867172	-30.18
9.0	mU to gU, finely decomposed organic	74.30	0.34	6.141	6.027	17.72675935	-30.26
12.0	mU, larger organic remains	59.21	0.225	3.725	3.303	14.68204233	-29.67
17.0	mU, larger organic remains	64.04	0.31	4.841	4.486	14.46965433	-29.21
22.0	mU, larger organic remains	70.54	0.333	6.48	5.888	17.68265186	-30.46
28.0	mU, larger organic remains	69.27	0.388	5.8	5.767	14.86434642	-30.09
30.0	mU, organic remains	64.23	0.389	5.878	5.686	14.61586854	-29.26
32.0	mU, organic remains	62.95	0.326	5.422	5.296	16.24415989	-29.32
34.0	mU, larger organic remains s decomposed	71.52	0.456	7.516	6.960	15.26260637	-30.20
38.0	mU, larger organic remains s decomposed	70.57	0.492	7.259	6.488	13.18697813	-29.57
40.0	mU, smaller organic remains	66.78	0.42	7.186	6.701	15.95480272	-30.00
45.0	mU, smaller organic remains	64.52	0.423	6.981	6.496	15.35661283	-29.41
50.0	mU, smaller organic remains	62.54	0.42	6.925	6.139	14.61612667	-28.64
53.0	moderately decomposed peat	78.21	1.144	21.156	19.882	17.37949148	-30.56
56.0	moderately decomposed peat	76.05	1.236	21.011	19.882	16.08587237	-30.83
Used for study							
Sample Depth [cm]	Short description	Water content [%]	TN [wt%]	TC [wt%]	TOC [wt%]	TOC/N	$\delta^{13}C$
1.0	mU, minerogenic	49.70	0,255	3.48	3.301	12.94521023	-28.38
3.0	mU to gU, finely decomposed organic	69.81	0,5	???	5.812	11.62400055	-29.26
6.0	mU to gU, finely decomposed organic	72.22	0,379	5.721	5.760	15.19867172	-30.18
9.0	mU to gU, finely decomposed organic	74.30	0,34	6.141	6.027	17.72675935	-30.26
AVERAGE		66.51	0,37	5.11	5.23	14.37	-29.52

Core ID	TOC	TC	TN	TOC/TN	13C
TLO18-12	6.8	8.11	0.56	12.19	-28.68
TLO18-15	5.72	6.11	0.38	14.87	-28.45
TLO18-26	11.337	11.487	0.674	11.337	-29.25
TLO18-27	4.28	4.35	0.27	15.64	-28.3
TLO18-36	8.527	8.954	0.622	13.766	-28.69
TLO18-48	6.94	7.04	0.42	16.61	-28.42
TLO18-52	4.66	4.82	0.3	15.59	-28.56
TLO18-54	4.438	4.5705	0.281	15.765	-28.48
TLO18-75	3.67	3.89	0.21	16.82	-27.16
TLO18-IKPO1	5.23	5.11	0.37	14.37	-29.52
LOW	0-4	0-4	0-0.3	11.3-12.5	(-29.01)-(-30)
MEDIUM	4.1-5.0	4.1-5.0	0.31-0.4	12.51-14.5	(-28.51)-(-29.00)
HIGH	5.1-8.0	5.1-8.0	0.41-0.6	14.51-15.5	(-28.31)-(-28.5)
VERY HIGH	8.1-12	8.1-12	0.6-0.7	15.51-17	(-27.16)-(-28.3)

Core ID	Size [km ²] 2018	Size [km ²] 1986	Dynamic [%]	Dynamic [km ²]	Depth [cm]	Ice lake type	Vegetation origin
TLO18-12	10.485	9.844	6.11349547	0.641	250	floating	Rather lacustrine
TLO18-15	10.018	9.572	4.451986424	0.446	220	floating	Rather terrestrial
TLO18-26	6.723	6.449	4.075561505	0.274	230	floating	Rather lacustrine
TLO18-27	3.758	4.028	-7.184672698	-0.270	64	grounded	Rather terrestrial
TLO18-36	5.877	5.535	5.819295559	0.342	200	floating	Rater Terrestrial
TLO18-48	3.987	3.809	4.464509656	0.178	160	grounded	terrestrial
TLO18-52	19.505	18.524	5.029479621	0.981	200	floating	Rather terrestrial
TLO18-54	17.404	16.727	3.889910365	0.677	100	grounded	Rather Terrestrial
TLO18-75	13.068	13.035	0.252525253	0.033	100	grounded	terrestrial
TLO18-IKPO1	0.689	0.652	5.370101597	0.037	350		Rather Terrestrial
Classification	0-4		±0.1-±2.9		0-100		
	4.1-8		±3-±4.4		101-200		
	8.1-12		±4.5-±5.9		201-250		
	12.1-18		±6-±7.2		251-400		

ERKLÄRUNG

Ich erkläre, dass ich die vorliegende Arbeit nicht für andere Prüfungen eingereicht, selbständig und nur unter Verwendung der angegebenen Literatur und Hilfsmittel angefertigt habe. Sämtliche fremde Quellen inklusive Internetquellen, Grafiken, Tabellen und Bilder, die ich unverändert oder abgewandelt wiedergegeben habe, habe ich als solche kenntlich gemacht. Mir ist bekannt, dass Verstöße gegen diese Grundsätze als Täuschungsversuch bzw. Täuschung geahndet werden.

Berlin, den

Linda Siegert